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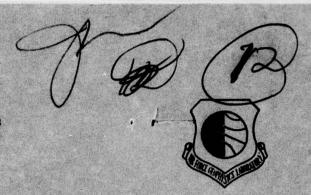
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LWIR (7-24µm) Measurements From the Launch of a Rocketborne Spectrometer Into an Aurora (1973)

JAMES W. ROGERS A.T. STAIR. Jr. NED B. WHEELER CLAIR L. WYATT DORAN J. BAKER

15 November 1976



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OPTICAL PHYSICS DIVISION PROJECT 2310

AIR FORCE GEOPHYSICS LABORATORY
HAMSCOM AFB, MASSACHUSETTS 01731

AIR FORCE SYSTEMS COMMAND, USAF



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FOR THE COMMANDER

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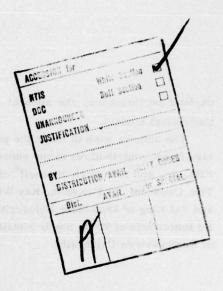
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This report documents in detail the data obtained along with the rocket and payload performance, the auroral conditions at launch and a comparison with theoretical models.

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Preface

The High Altitude Effects Simulation (HAES) Program sponsored by the Defense Nuclear Agency since the early 1970 time period, comprises several groupings of separate, but interrelated technical activities, for example, ICECAP (Infrared Chemistry Experiments - Coordinated Auroral Program). Each of the latter has the common objective of providing information, ascertained as essential for the development and validation of predictive computer codes designed for use with high priority DoD radar, communications, and optical defensive systems.

Since the inception of the HAES Program, significant achievements and results have been described in reports published by DNA, participating service laboratories, and supportive organizations. In order to provide greater visibility for such information and enhance its timely applications, significant reports published since early calendar 1974 have been identified with an assigned HAES serial number and the appropriate activity acronym (for example, ICECAP) as part of the report title. A complete and current bibliography of all HAES reports issued prior to and subsequent to HAES Report No. 1, dated 5 February 1974 entitled "Rocket Launch of an SWIR Spectrometer into an Aurora (ICECAP 72)", AFCRL Environmental Research Paper No. 466, is maintained and available on request from DASIAC, DoD Nuclear Information and Analysis Center, 816 State Street, Santa Barbara, California 93102, Telephone (805) 965-0551.

This report, the fifty-first in the HAES series, presents the data obtained from the launch of a liquid-helium-cooled, long-wavelength infrared spectrometer into an aurora. While preliminary data have been presented previously (see references 1 and 24), this report documents in detail the data obtained along with the rocket and

payload performance, the auroral conditions at launch and a comparison with theoretical models.

The authors wish to thank the personnel from various organizations who made significant contributions to the success of this experiment including: Dr. H.C. Fitz, Jr. and Mr. Herb Mitchell* of DNA; Neil Brown of The University of Alaska; Tom Condron, Ed McKenna, Ray Wilton and Floyd Cook of AFGL; Gary Frodsham and Val King of Utah State University; Ed Allen of Space Data Corporation; Ed Butterfield of White Sands Missile Range and Larry O'Connor and Dick Morin of Northeastern University.

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^{*}Presently with R&D Associates

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LWIR (7-24 µm) Measurements From the Launch of a Rocketborne Spectrometer Into an Aurora (1973)

1. INTRODUCTION

A liquid-helium-cooled, long-wavelength infrared (LWIR) spectrometer was successfully launched by the Air Force Geophysics Laboratory (AFGL)* onboard a Black Brant VC rocket (A18.006-2) on 22 March 1973 from the University of Alaska's Poker Flat Research Range at Chatanika, Alaska. This flight was part of the Defense Nuclear Agency (DNA) ICECAP 73 Program. The spectrometer, which employs a circular-variable filter (CVF), was developed by AFGL and Utah State University (USU) and has provided the first measurements of the altitude profile of the infrared spectrum of the upper-atmospheric emissions between 7 and 24 μm .

During an IBC II aurora, data were obtained on the 15- μ m carbon dioxide (ν_2) emission from 65 to 150 km and on the 9.6- μ m ozone (ν_3) emission between 45 and 100 km as reported by Stair, et al. ¹ Emission data on the long-wavelength wing of the 6.3- μ m water band between 6.7 and 7.6 μ m were also obtained between 50 and 75 km. Above 100 km, significant unidentified emission was observed at 9.3 μ m with weaker features at 6.9, 7.3, 8.0, 11.1, and 12.3 μ m. (Received for publication 12 November 1976)

Formerly the Air Force Cambridge Research Laboratories.

Stair, A.T., Jr., Ulwick, J.C., Baker, K.D. and Baker, D.J. (1975) Rocket-borne observations of atmospheric infrared emissions in the auroral region, Atmospheres of Earth and Planets, edited by B.M. McCormac (D. Reidel, Dordrecht-Holland), pp 335-346.

The calibration particulars and an analysis of the spectrometer performance along with the methods of processing the data have been presented in separate reports by Condron² and Rogers.³ This report documents in detail the data obtained along with the rocket and payload performance, the auroral conditions at launch and a comparison with theoretical models.

2. PAYLOAD CONFIGURATION

In addition to the liquid-helium-cooled LWIR CVF spectrometer, the rocket payload was instrumented with diagnostic sensors to provide information on local auroral conditions plus rocket and payload performance. These included a 3914 Å photometer, a light-scattering dust detector, mechanical vibration spectrometers, an aspect gyro, and a spin magnetometer.

The rocket was assembled with a "clamshell" split-nosecone which could be opened during ascent at a predetermined altitude above which the LWIR sensor would not seriously cryo-pump. The clamshell could then be reclosed during descent. The payload was outfitted with a recovery system employing a parachute and beacon. A dust detector was provided to indicate the presence of local particulate matter which might pass through the field of view of the spectrometer. To minimize possible dust contamination problems, all portions of the payload were sealed, and both the sensor and the payload were assembled in a clean room. The final payload was then enclosed in a large plastic bag until launch.

A summary of the field requirements, along with the specifications of the Black Brant VC rocket can be found in a field requirements report by Space Data Corporation.

2.1 LWIR CVF Spectrometer Description

The spectrometer flown was Model HS-1B-1B developed by USU and is described as follows. The spectral region from 7 to $24\,\mu\mathrm{m}$ is scanned at a rate of two scans per second with a full-angle field of view of 2 deg $(9.6\times10^{-4}~\mathrm{sr})$. The entire optical subsection is cooled to liquid-helium temperature, including an arsenic-doped silicon

Condron, T. P. (1973) Calibration of a Liquid Helium Cooled CVF Radiometer in a Warm Environment, AFCRL Instrumentation Papers, No. 199, AFCRL-TR-73-0480.

Rogers, J.W. (1975) Instrumentation Analysis and Data Processing for Rocketborne LWIR Spectrometers (With Application to Rocket A18.006-2 of 22 March 1973), Environmental Research Papers, No. 539, AFCRL-TR-75-0535, HAES Rpt. No. 23.

ICECAP '73 Field Requirements and Procedures (1973) SDC TM-620, Revision D, Space Data Corporation, Phoenix, Arizona, 6 March 1973.

detector which is thereby operated with an effective zero radiation background. A key element of the spectrometer is the Optical Coating Laboratory, Inc. (OCLI) rotating variable interference filter. The spectral resolution of the filter was measured at AFGL by Condron and the best estimate of the half-bandwidth for the flight- configured spectrometer was 2.63 percent for the short-wavelength filter half and 2.92 percent for the long-wavelength filter half. This circular variable filter element has led to this class of instruments being referred to as CVF spectrometers. A dynamic range of 3×10^4 is achieved by providing four data channels with different gains. The output data from the spectrometer along with the diagnostic information were transmitted to a ground telemetry site during the flight.

Many technical breakthroughs had to be achieved in order to assure satisfactory measurements. Most of these problems dealt with the cryogenics and the stringent requirements placed on the electronics and optics to perform remotely in a severe environment. A cutaway view of the spectrometer is shown in Figure 1. Only general information on the spectrometer will be given in this report while more specific details can be found in reports by Wyatt⁵⁻⁷ and Stair, et al. 8

The CVF used is composed of two 180-deg wedge filter segments mounted together to form a disc. The short-wavelength half (rotational angles 0 to 180 deg) covers wavelengths from 6.5 to 13 μ m. The long-wavelength half (angles from 180 to 360 deg) covers the range from 12.5 to 25 μ m. Small portions of these spectral regions are not usable because opaque metal joining strips are placed over the interfaces between the two filter halves at 0 deg and at 180 deg. Each time the metal strip at 0 deg passes in front of the detector, a feedback circuit cancels the residual voltage caused by any amplifier dc drift. The voltage pulse which triggers this dc reset is monitored and outputted on the calibration reference channel of the spectrometer.

The rotational position of the CVF is monitored during each scan by a coded chopper that is attached to the same shaft as the CVF. This encoder is located at a distance of twenty centimeters behind the filter. The chopper periodically interrupts radiation from an incandescent source that is incident upon a silicon detector.

^{5.} Wyatt, C.L. (1971) Infrared Helium-Cooled Circular-Variable Spectrometer
Model HS-1, Final Report, AFCRL-71-0340, Contract No. F19628-67-C-0340,
Utah State University, Logan, Utah, September 1971.

Wyatt, C. L. and Baker, D. J. (1975) Development of a Liquid-Helium Cooled Rocketborne Spectrometer, Scientific Report No. 2, AFCRL-TR-75-0164, Contract No. F19628-73-C-0048, Utah State University, Logan, Utah, February 1975.

Wyatt, C. L. (1975) Infrared Spectrometer; Liquid-helium-cooled rocketborne circular-variable filter, <u>Applied Optics</u>, 14, No. 12:3086-3091.

^{8.} Stair, A.T., Jr., Wheeler, N.B., Baker, D.J., and Wyatt, C.L. (1973)
Cryogenic IR spectrometer for rocketborne measurements, NEREM 73
Record, IEEE Catalog No. 73, CHO-841-7 NEREM, pp 80-89.

The end result is a series of 10 coded pulses produced during each filter rotation which are then outputted on the position reference channel. In principle, the rotational angle of the filter at each instant, and, therefore, the wavelength of peak transmission at that time, can be uniquely ascertained by the train of pulses produced by the chopper.

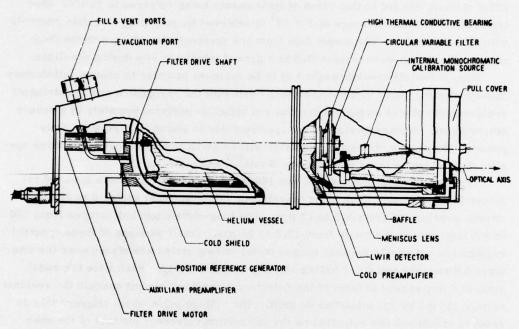


Figure 1. Cutaway View of the Liquid-Helium Cooled CVF Spectrometer Used for IR Atmospheric Emissions

The spectrometer also contains an internal calibration source which is activated every tenth scan. This source consists of a gallium-arsenide emitter which scatters unfiltered light onto the detector. Due to its temperature sensitivity, the internal source is insufficiently stable to provide an absolute inflight calibration of the spectrometer; however, the emitter can be used as a measure of the spectrometer performance. The internal calibration pulse voltage is monitored on the calibration reference channel output.

The final flight calibration of the spectrometer's absolute spectral response was performed at AFGL and is contained in a report by Condron. ² The spectral response in W cm⁻²sr⁻¹ μ m⁻¹/V of the highest gain channel is shown in Figure 2 for both halves of the filter and is accurate to within a factor of two.

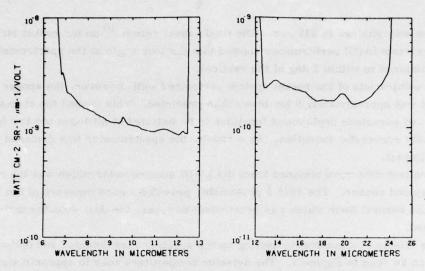


Figure 2. Final Flight Calibration of the CVF Spectrometer

2.2 Photometer Description

A 3914 Å photometer was included in the payload to provide a measure of the auroral energy deposition. The photometer field of view was coaligned with the LWIR spectrometer and had a value of 10.3 deg full angle (0.0254 sr). The transmittance of the interference filter peaked at 3907 Å with a spectral bandwidth of 17 Å. The unattenuated dynamic response of the photometer ranged from 0.2 to 500 Rayleigh/Å.

The photometer operated directly coupled in an analog mode with a dual-range linear channel. A "dogleg" gain change of 10 was used to extend the dynamic range over the single linear telemetry channel that was available.

3. ROCKET AND PAYLOAD PERFORMANCE

The Black Brant VC rocket (A18.006-2) with the LWIR CVF spectrometer payload was launched at 1212:57 hours UT on 22 March 1973. The payload was integrated at AFGL under the direction of E. McKenna with R. Wilton of AFGL responsible for telemetry and tracking. Trajectory data are contained in a report prepared by the Analysis and Simulation Branch of the AFGL Computation Center. The altitudes reported are accurate to within 0.5 km and indicate that an apogee of

^{9.} Trajectory Report, Rocket No. A18-006-2 (1973) Analysis and Simulation Branch, AFCRL Computation Center.

185.4 km was attained in 217 sec. The final aspect report ¹⁰ on the rocket attitude control system (ACS) performance showed that the look angle of the spectrometer was maintained to within 2 deg of the vertical.

All components of the rocket system performed well; however, the apogee obtained was approximately 9 km lower than predicted. This caused the clamshell closing and parachute deployment functions to be initiated at altitudes too low for completely successful execution. As a result, the spectrometer was damaged upon ground impact.

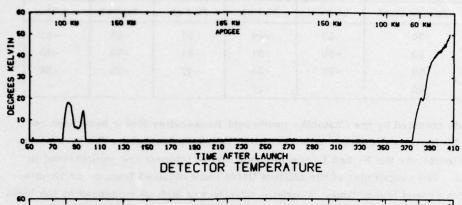
Excellent data were obtained from the LWIR spectrometer which was the primary payload sensor. The 3914 Å photometer provided a good measure of the intensity of the auroral form which was penetrated; however, the dust detector malfunctioned.

The effects of atmospheric heating on the spectrometer detector and optical baffle can be seen in Figure 3. The detector temperature rose to approximately 18° K immediately after the clamshell was opened at 70.5 sec (86 km), but returned to the temperature of liquid helium by 97 sec (119 km). During payload descent, the temperature started to rise at 375 sec (69 km) and was 35° K during the last usable data scan at 390 sec (45 km). The baffle temperature also rose after the clamshell opened and peaked at 50° K around 120 sec (142 km). This was well below the upper limit that would have affected the spectrometer performance. The baffle cooled off throughout the flight until around 365 sec (83.5 km). This cooling is attributable to the radiation of the heat energy which had built up during the clamshell opening shock. During the last usable data scan, the baffle was at 45° K.

4. LAUNCH CONDITIONS

A post-breakup auroral glow had persisted at the IBC Class II level for nearly 10 min prior to the penetration of the rocket into the region. At the time and location of penetration, the zenith radiance was 16.5 kR for the 5577 Å auroral green line as measured by The University of Alaska from Fort Yukon. The attitude of the optical axis of the spectrometer was inclined at an angle of 12.8 deg north relative to the magnetic field lines above Poker Flat. The aurora remained relatively stable during the rocket flight, but small scale changes occurred over the whole sky. The phenomenological situation appeared to be that the rocket was launched during an intense auroral substorm shortly after the poleward expansion had occurred.

^{10.} Aspect Report, Rocket No. A18-006-2 (1974) Aspect Report Number 4531-9, Analysis and Simulation Branch, AFCRL Computation Center.



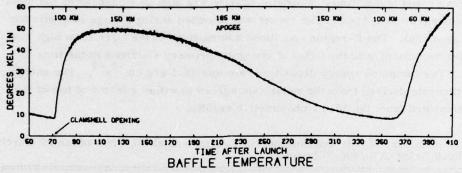


Figure 3. Spectrometer Detector and Baffle Temperatures During Flight

Additional aeronomic information pertinent to this experiment, including atmospheric meteorological data, ground station magnetometer, all-sky camera and photometric measurements can be found in the report from the Geophysical Institute of The University of Alaska by Romick. ¹¹ The reported temperature profiles obtained from sounding rocket data are shown in Table 1. These temperatures help describe the atmospheric environment and are very important in analyzing the flight data. The characterization of the aurora during the flight has also been reported by Kofsky, et al¹² of Photometrics, Inc.

^{11.} Romick, G.J. (1975) Report of the Geophysical Description and Available Data
Associated with Rocket PF-BB-53, Scientific Rpt No. 2, AFCRL-TR-750040, Contract No. F19628-74-C-0188, Geophysical Institute, University
of Alaska, January 1975.

Kofsky, I. L., Meriwether, J. W., Schroeder, J. W., and Sluder, R. B. (1974)
 Data Reduction and Auroral Characterizations of ICECAP, HAES Report
 No. 4, Final Report, DNA 3511F, Contract No. DNA001-73-C-0027,
 Photometrics, Inc., Lexington, Mass., September 1974.

Table 1. Poker Flat Temperatures (°C) From Rocket Sounding Data (1973)

Altitude (km)	14 Feb	21 Feb	22 Feb	24 Feb	1 Mar
30	-55	-54	-57	-55	-59
40	-50	-51	-51	-52	-45
50	-39	-33	-37	-35	-24
60		-17			

Data obtained by the Chatanika Incoherent Backscatter Radar have been reported by Baron and Chang 13 of the Stanford Research Institute. The parameters that characterize the F- and E-region conditions of interest are summarized in Table 2. The magnitudes of the entities listed were obtained from 5- or 10-min averages around launch time. Magnetic activity was high as evidenced by the local K-indices shown in Table 2. The rocket was launched during a large magnetic bay (-1500 gammas). The E-region conditions summarized in the table show high peak densities associated with the influx of energetic primary electrons at the time of launch. The computed energy deposition rate was 16.1 erg cm⁻²s⁻¹. The energy deposition rate derived from the radar data agreed to within a factor of two of the value computed from the 3914 Å photometric results.

Table 2. Summary of Selected Ionospheric Parameters at the Time of the 22 March 1973 Black Brant (A18, 006-2) Launch

Local 3-hr K-index	8			
F region				
N _e (300 km)	6 × 10 ⁴	el/cm ³		
T _e (230 km)	1350	°K		
T _i (230 km)	900	•K		
E region				
N _e (max)	6.6×10^5	el/cm ³		
he (max)	90	km		
Particle energy	50	keV		
Number flux, Φ	2 × 10 ⁸	cm ⁻² s ⁻¹		
Electric fields:				
GM North	-16	mV/m		
GM East	- 9	mV/m		
Joule heating	11	ergs/cm ⁻² s ⁻¹		
Particle energy flux:	Photos I say to Percoll			
Radar	16.1	ergs/cm ⁻² s ⁻¹		
Optics	16.4	ergs/cm ⁻² s ⁻¹		
Joule heating/particle energy flux	0.67	Profit Line Contact		

^{13.} Baron, M. and Chang, N.J. (1975) ICECAP '73A - Chatanika Radar Results, HAES Report No. 15, Topical Report, DNA 3531T, Contract No. DNA 001-74-C-0167, Stanford Research Institute, Menlo Park, California, April 1975.

5. EXPERIMENTAL RESULTS

The processing of the LWIR spectrometer data proceeded in a step-by-step manner to provide accurate final data complete with error limits. This also made it possible to assay the system performance. An analysis, with the determination of which data scans were actually usable, has been presented in a separate report by Rogers. After the payload clamshell was opened at 86.4 km, 680 scans (revolutions of the CVF) were obtained of which a total of 488 were considered usable. Scans were discarded when all channels were saturated, during internal calibration, when gamma ray disturbances occurred, and during the time of passage of an unidentified object through the spectrometer's field of view.

5.1 N2 (1N) Altitude Profile

The measured altitude profile of the overhead radiance at 3914 Å due to the first negative band ($B^2 \Sigma_u^+ \to \chi^2 \tau_g^+$) of N_2^+ (0,0) in the near ultraviolet is given in Figure 4 as a function of rocket altitude. The half-power bandwidth of the photometer filter was 17 Å with a field of view of 10.3 deg full angle. On ascent, the overhead photon radiance decreased from about 5 kR at 100 km to 0.8 kR at 170 km as the rocket passed through the auroral form. Overhead at 185 km, the 3914 Å radiance was less than 140 R. During descent, the intensity below 140 km was consistently lower than on the upleg trajectory.

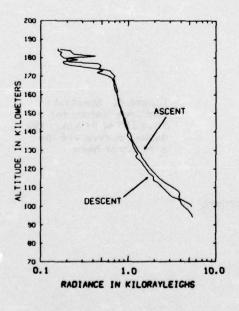


Figure 4. The Altitude Profile or Overhead Radiance at 3914 Å as Measured Onboard the Rocket

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5.2 LWIR CVF Infrared Spectra

Figures 5 and 6 show typical examples of individual scans obtained with the dashed lines representing the ± 30 error limits. Scan 477 (Figure 5) was taken 219.4 sec after launch near apogee at 185 km and scan 755 (Figure 6) was at 357.1 sec (94 km) during the downleg portion of the trajectory. At the higher altitudes, where the signals are weak and varying slowly, the data scans can be averaged to reduce the random noise at each wavelength. The resulting gain in signal-to-noise from this process is shown at apogee in Figure 7 where the spectral radiance values have been averaged over 34 scans. Appendix A contains data scans obtained throughout the entire flight. Averaging has been performed as indicated in the figure captions.

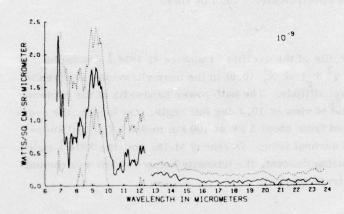


Figure 5. Spectral Radiance Values for Scan 477 at 185 km. The dashed curves are the ± 30 error bars

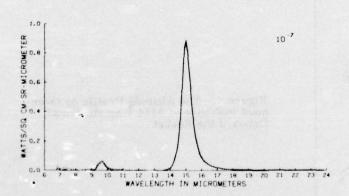


Figure 6. Spectral Radiance Values for Scan 755 at 94 km. The dashed curves are the ± 3σ error bars

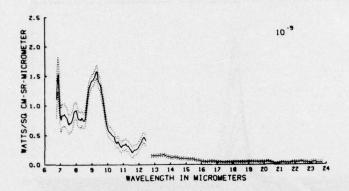


Figure 7. Spectral Radiance Values at Apogee Obtained by Averaging Over 34 Scans. The dashed curves are the ± 30 error bars

Graphic portrayals of the spectral intensities measured during this experiment are given by three-dimensional presentations as shown in Figures 8, 9, and 10. Due to the dynamic range of the spectrometer some of the noise baseline features are lost in these plots. Figure 8 shows a portion of the ascent data between 100 and 185 km where the scans have been averaged over 2.5-km increments. The decay of the 15- μ m CO₂ emission is clearly seen while the 9.6- μ m O₃ emission is only shown near 100 km. At the higher altitudes, the unidentified emissions are evident at the shorter wavelengths. Figure 9 presents a similar view of the descent data from 185 to 100 km. Every scan from 100 to 50 km during descent is shown in Figure 10. The clipping of the 15- μ m CO₂ at the lower altitudes is caused by amplifier saturation.

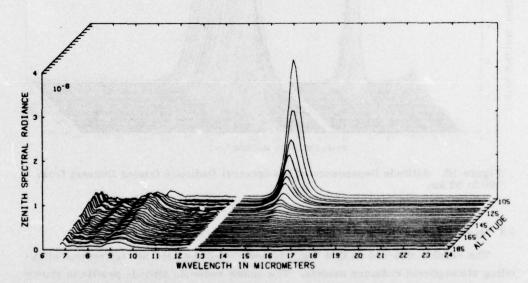


Figure 8. Altitude Dependence of the Spectral Radiance During Ascent from 100 to 185 km. Spectral scans are averages over 2.5 km increments

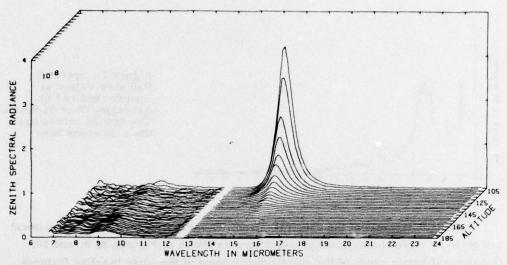


Figure 9. Altitude Dependence of the Spectral Radiance During Descent from 185 to 100 km. Spectral scans are averages over 2.5 km increments

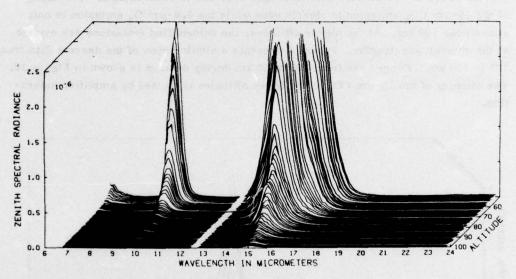


Figure 10. Altitude Dependence of the Spectral Radiance During Descent from 100 to $50\ km$

5.3 O3 and CO2 Zenith Radiance Profiles

The results obtained on this flight demonstrated the need for refinement of existing atmospheric radiance models. The zenith radiance altitude profile is shown in Figure 11 for the 9.6- μ m O_3 and in Figure 12 for the 15- μ m CO_2 . The dashed

lines represent the ± 30 error bars of the spectrometer data. Additional errors in the altitude profiles are introduced in altitude determination (0.5 km) and spectrometer calibration (a factor of 2). The larger spread between the data and the error bars at the higher altitudes indicates that the spectrometer's lower sensitivity level is being reached. At altitudes below 75 km, the measured ozone

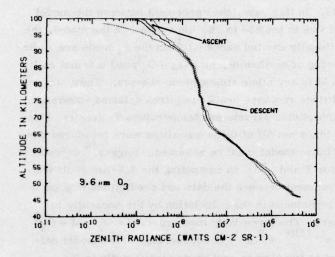


Figure 11. Zenith Radiance Altitude Profile at 9.6 μ m Measured With the LWIR Spectrometer. The dashed curves are the \pm 3 σ error bars

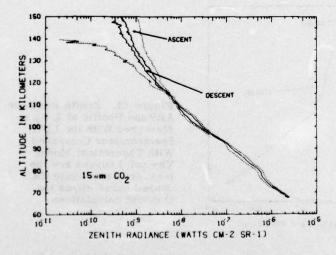


Figure 12. Zenith Radiance Altitude Profile at 15 μ m Measured With the LWIR Spectrometer. The dashed curves are the \pm 3 σ error bars

emission values are in reasonable agreement with the most accurate models at the time of the flight (Corbin et al. 14 and Degges 15), but are two orders of magnitude larger at higher altitudes. The measured ${\rm CO_2}$ emission data also agrees with the models of Corbin et al. 14 and Degges 15 only at the lower altitudes. The disagreement is a factor of 10 at 85 km and 2 orders of magnitude at 150 km.

Comparison between the measurements and a more recently calculated radiance profile by Degges 16 using the ozone profile of Roble and Hays 17 gives a much better agreement as shown in Figure 13. In this case, the agreement between the model and the data is within a factor of five in the 60- to 100-km range. In the model, the principal mechanisms for vibrationally excited ozone (O_3^{-1}) in the ν_3 mode are radiation transport; resonance scattering of earthshine, $h\nu + O_3 \rightarrow O_3^{\dagger}$; and thermal collisions, $O_3 + M \rightarrow O_3^{\dagger} + M$, where M is any other atmospheric species. Thus, accurate modelling of the $9.6-\mu m$ altitude radiance profile requires detailed knowledge of the altitude profiles of the atmospheric parameters temperature T; density {M;} and ozone concentration {O3} . Since not all of these quantities were measured during this rocket flight, an atmospheric model must be assumed. Degges 16 describes the atmospheric model he used for T and $\{M\}$ in computing the 9.6- μ m radiance profile shown in Figure 13. Agreement between the data and predictions is good if one considers the uncertainties introduced in the calculation by the necessity to model the atmospheric parameters. The three-body mechanism $O + O_2 + M \rightarrow O_3^{\frac{1}{3}} + M$ (with a rate of 3.2 × 10⁻³⁵ e^{1.7/Rt} cm⁶ s⁻¹) is not included in the model calculations shown in Figure 13, but is too slow to contribute significantly to the

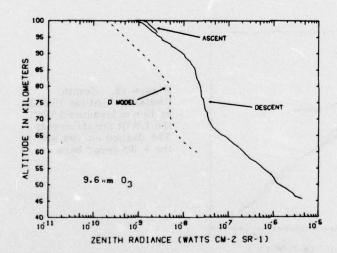


Figure 13. Zenith Radiance Altitude Profile at 9.6 μ m Measured With the LWIR Spectrometer Compared With Theoretical Model. The solid curves are the measurements while the dashed curve shows the D model calculations

The references cited on the next several pages are too numerous to list here; see References on page 31.

emission above 80 km. However, it does become important below 80 km and its inclusion would place the theoretical model and experimental measurement in closer agreement in this altitude region.

Two new model calculations for the 15 μ m CO $_2$ radiance profiles are compared with the measured results in Figure 14. These are the Degges 16 model (D) and the results (KJ) of an independent calculation by James and Kumer 18 and Kumer 19 . In these models, the principal mechanisms for producing CO $_2$ in the vibrationally excited ν_2 state (CO $_2^{\frac{1}{4}}$) are radiation transport $h\nu + CO_2 \rightarrow CO_2^{\frac{1}{4}}$ and thermal collisions $CO_2 + M \rightarrow CO_2^{\frac{1}{4}} + M$. Again, accurate modelling requires a knowledge of the altitude profile of the atmospheric parameters, T, $\{M\}$, and $\{CO_2\}$. These were modelled in the two cases as described by Degges 16 and Kumer. Differences in the D and KJ calculations largely reflect the differences in the atmospheric models employed. Also, the CO_2 mixing ratio in the D model remains constant at essentially 3.1×10^{-4} above 100 km whereas the KJ model utilizes the smaller CO_2 mixing ratio above 100 km which was introduced by James and Kumer. Neither model accounts for the enhanced scattering of earthshine in the region above 100 km which should result from the CO_2 band and line broadening due to the increased temperature in the thermosphere. These models are in excellent agreement at the lower altitudes with the measurements, but exhibit discrepancies increasing with altitude.

Indirect evidence supports the claim that the experimental measurements are accurate to within a factor of two. First of all, upon extrapolation, the rocket data joins smoothly with the balloon measurements of Murcray et al. ²¹ (Stair et al. ²⁴). Secondly, satellite measurements tend to support the low altitude data (Conrath

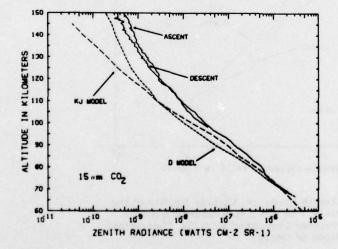


Figure 14. Zenith Radiance Altitude Profile at 15 μ m Measured With the LWIR Spectrometer Compared With Theoretical Models. The solid curves are the measurements while the dashed curves show the model calculations of Kumer and James (KJ) and Degges (D)

et al. ²²; Barnett et al. ²³ Finally, preliminary analysis of the 1974 LWIR CVF rocket measurements under ambient conditions gives approximately the same intensity as under auroral conditions (Stair et al. ²⁴). This was expected since neither band is predicted to be enhanced by aurora to any significant extent (Bishop et al. ²⁵).

5.4 H₂O Zenith Radiance Profile

Apparent high altitude emission from the long-wavelength wing of the 6.3- μ m water band is evident in Figure 10. The resulting zenith radiance altitude profile obtained by integrating the data from 6.7 to 7.6 μ m is shown in Figure 15. Also shown by the dashed curve is a theoretical calculation for the water radiance profile in this wavelength region using the model of Degges ¹⁶ and the fact that 20 percent of the 6.3- μ m water band emits between 6.7 and 7.6 μ m. The calculation used a standard arctic atmosphere with 5 ppm H_2O and temperatures from Table 1. The primary excitation processes for H_2O at these altitudes are earthshine $h\nu + H_2O \rightarrow H_2O^{\frac{1}{2}}$, and the transfer of vibrational energy through thermal collision $H_2O + M \rightarrow H_2O^{\frac{1}{2}} + M$. De-excitation of vibrationally excited H_2O is by radiation and collisional deactivation. Collisional excitation and de-excitation is mainly due to collisions with molecular oxygen and to a lesser degree with molecular nitrogen. The resulting agreement between the calculated and measured data supports the contention that the emission is indeed from the 6.3- μ m water band.

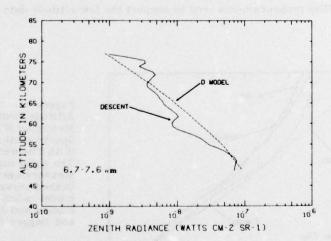


Figure 15. Zenith Radiance Altitude Profile of the Emission from 6.7 to 7.6 μ m Measured With the LWIR Spectrometer Compared With the Theoretical $\rm H_2O$ Calculations of Degges

5.5 Unidentified Emissions

Above 100 km, significant unidentified emission was observed at 9.3 μ m with weaker features at 6.9, 7.3, 8.0, 11.1, and 12.3 μ m. All these emission bands can be seen in Figure 6 with the feature at 9.3 μ m being a particularly strong emitter. Figure 16 is the zenith altitude profile of the 9.3 μ m feature which shows a marked asymmetry between ascent and descent. The observed radiance values are plotted vs time in Figure 17 where it becomes obvious that, after an initial rise, the intensity decreases throughout the remainder of the flight. Figure 18 shows the time variations of the weaker unidentified emissions which were observed above 100 km. It is seen that all these features have a similar time dependence as the larger feature at 9.3 μ m. The horizontal velocity component of the payload is about 0.55 km/sec and during the time shown in these plots, the horizontal range from the launch point increases approximately 150 km.

Further examination of Figure 6 and the scans above 100 km in Appendix A, shows that in addition to the spectral features mentioned there is a background continuum which increases towards the shorter wavelengths. The decrease in the intensities of the emissions shown in Figures 17 and 18 is caused by both the unidentified bands and continuum background. For the zenith radiance profile between 8.5 and 10.9 μ m shown in Figure 18, the continuum accounts for approximately 43 percent of the total emission up to 300 sec when it drops off to less than 20 percent.

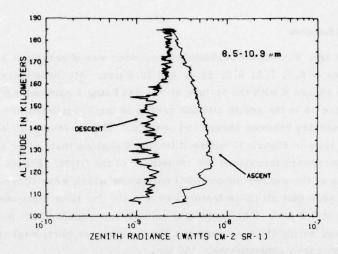


Figure 16. Zenith Radiance Altitude Profile of the Strong Unidentified Emission at 9.3 $\mu\,m$

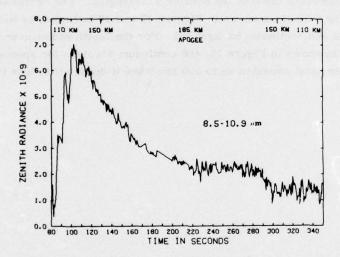


Figure 17. Zenith Radiance of the Unidentified Emission Feature at 9.3 $\mu\,m$ as a Function of Time After Launch

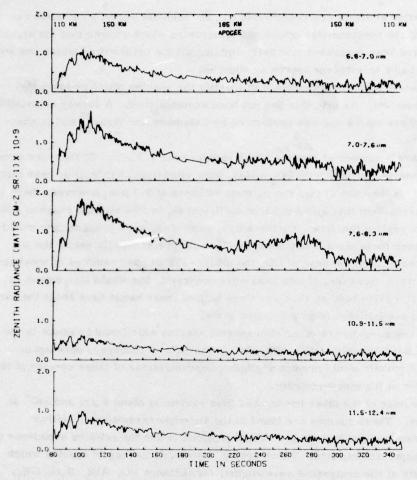


Figure 18. Zenith Radiance Values of Weaker Unidentified Emissions as Functions of Time After Launch

There are a number of possible explanations regarding these emissions. The first is that a yet unidentified atmospheric radiation was in the instruments's field-of-view during ascent, but due to its temporal and/or spatial behavior moved out of the field-of-view during descent. Auroral related emissions, for example, could have such profiles. Secondly, it is possible that these spectral features are due to exo-atmospheric sources such as dust clouds, stars, galaxies, and so forth. A third possibility is that the emissions are due to contaminants introduced by the measuring system or from the rocket exhaust which were carried along during the payload ascent and diffused at the higher altitudes. Finally, the possibility that

instrumentation artifacts are responsible can be ruled out due to a complete reevaluation of the spectrometer optics and electronics which showed that the signals could not have been generated internally and that all the unidentified emissions were indeed caused by an external source or sources.

It is clearly of prime importance to identify the species which produce the observed radiation. As yet, this has not been accomplished. A survey of possible sources of these emissions was performed by Caledonia and Wray 26 and is summarized here.

There are a number of species which radiate near 9.3 μ m. Of the more prominent atmospheric species, both CO₂ and O₃ have vibrational bands at or near this wavelength. In the case of O₃, the ν_1 state radiates at 9.3 μ m; however, there is no known mechanism that would produce sufficient O₃ in this altitude range of 100 to 185 km to explain the data. Furthermore, even if O₃ were present, one would have to further postulate a mechanism which would preferentially excite the ν_1 state over the ν_3 . In the case of CO₂ the (001) \rightarrow (02°0) band radiates at approximately 9.4 μ m. However, if this band were observed, one would also expect to see the (001) \rightarrow (100) band at 10.4 μ m since both of these bands have about the same lifetime and are initiated from a common state.

There are a number of other atmospheric species which would radiate in the region of 9.3 μ m notably, CHO, HNO, HO₂, and H₃O⁺. However, once again atmospheric models would predict negligible concentrations of these species at the high altitudes of the measurements.

For the case of the other bands, SiO does radiate at about $8~\mu\,\mathrm{m}$ and SiO^+ at about $12~\mu\,\mathrm{m}$. These species are found in the atmosphere; however, neither measurements nor predictions are available which define the relative abundance of SiO in the upper atmosphere. There are, of course, a number of species which would radiate at the designated wavelengths, for instance PO, AlN, B₂O, CH₂, COS, FeO, WO, PO₂, and so forth. Such species are not found in the natural atmosphere but might arise through instrumentation contamination.

A column density of the radiating species can be estimated from the data at 9.3 μ m. For a band radiance of 3 \times 10⁻⁹ W cm⁻² sr⁻¹ and a radiative lifetime of 10⁻² sec, the required excited state column density would be 1.5 \times 10¹⁰ part-cm⁻².

The total species column density cannot be specified unless the excitation source is defined. For earthshine excitations and radiative decay as the only loss mechanism, a column density of 10^{13} part-cm⁻² would be required at a base altitude of 185 km. For comparative purposes, the column density of N_2 between 180 to 300 km

^{26.} Caledonia, G. E. and Wray, K. L. (1974) Investigation of Auroral Phenomena and Light Scattering Models, Final Report, PSI-TR-5, Subcontract No. 129234B to Utah State University, Contract No. F19628-73-C-0048, Physical Sciences, Inc., Wavefield, Mass.

is approximately 2×10^{16} part-cm⁻². Based upon these numbers, it would appear that earthshine would be an unlikely excitation source for the observed radiation.

It is also possible to consider that the unknown radiation was collisionally excited. Assuming again that radiative decay is the only loss mechanism, a collisional rate constant of 10^{-10} cm³ sec⁻¹ and a collision partner number density of 10^{10} part-cm⁻³ (approximate atmospheric number density at 185 km) a lower bound requirement for the column density of 1.5×10^{12} part-cm⁻² is obtained. This bound is very crude and if collisions were the excitation source, most probably the column density of radiating species required to produce the observed radiation would be much larger.

With regard to the estimated excited state column density, it should be mentioned that if the observed radiation were due to instrument contamination, the molecular impurity levels would have to be relatively large. For example, for a radiative lifetime of 10^{-2} sec and a contamination length of one meter, the contaminant density would have to be larger than 1.5×10^8 part-cm⁻³ and this should be a strong lower bound.

There are a number of other possible radiation sources which cannot be ruled out. It is conceivable, for instance, that the observed bands could arise from electronic transitions of atmospheric species. For example, little is known about the high lying Rydberg states of N_2 . High level Rydberg state transitions might well occur in the $10-\mu$ m region and highly nonequilibrium electronic state populations could be generated during an auroral event.

Another possibility would be particulate emission where the emissivity of the particulate material was highly structured so as to give the appearance of band radiation. For purposes of example, consider spherical particles at a temperature of 200°K and unity emissivity. For a very large particle, say $r=10~\mu$ m, 9 part-cm⁻² would be sufficient to explain the data. On the other hand, for 1- μ m radius particles, a column density of order 1.5 × 10³ part-cm⁻² would be required.

On the surface, one would not expect particulate radiation to represent a viable source for the observed radiation, in particular because of the highly structured emissivity that would be required. Nonetheless, this source cannot be ruled out completely.

The possibility that the observed radiation arises from exo-atmospheric sources must be considered. For example, the zodiacal light has been postulated to be made up of particles of "dirty quartz", and SiO and SiO₂ have spectral bands in the wavelength regions of the unidentified emitters. The observed variation in intensity vs time (or apparent observation position due to payload velocity) appears to argue against this explanation. Known infrared celestial sources such as stars, galaxies,

and so forth, were eliminated as candidate sources by comparison with AFCRL Infrared Sky Survey. 27

Another potential source would be the vibrational/rotational bands of interstellar molecules. The radiance of these bands is not readily amenable to prediction and satellite measurements would be of great value in determining their importance. Many interstellar molecules have been identified to date, for the most part through the identification of rotational transitions in the mm to cm wavelength region. A number of isotopic variations of these molecules has also been detected.

These molecules are typically found in interstellar molecular clouds. With the exception of H_2 , which is by far the most abundant interstellar molecule, estimated line densities of these molecules are of order 3×10^{12} to 3×10^{16} part-cm⁻². The estimated CO line density falls at the upper end of this range while that for most other molecules falls at the lower. The typical kinetic temperatures of interstellar clouds are predicted to lie between approximately 20 to 125° K while observed rotational state temperatures have typically been found to fall between 4 to 45° K. It should be emphasized that the rotational and vibrational populations of these molecules should be severely collision limited.

Evidence of high altitude emissions in this wavelength region has been observed on two subsequent flights. Preliminary analysis of the 1974 LWIR CVF (with no auroral activity) shows evidence of an emission band around 9.3 μ m with a slightly different band shape; however, the background continuum is not evident. Preliminary analysis 28 of a rocketborne LWIR radiometer flown during ICECAP 75 into an intense IBC III auroral breakup also indicates that above 100 km there are emissions in the 7 to 12 μ m region at levels higher than predicted for ozone; however, the time history of the radiation showed an increasing value throughout the flight which is just the opposite of the 1973 LWIR CVF spectrometer. The radiometric measurements would not, however, rule out emissions from contaminant exhaust gases as it used a different boost rocket and the trajectory of any unwanted radiators could be quite different.

To resolve whether the emissions are due to possible instrumentation contaminants or exhaust gases or if they are indeed emissions from unidentified atmospheric or exo-atmospheric radiators will require additional experiments, preferably with higher spectral resolution for positive identification and using a "clean" booster or observation platform.

^{27.} Walker, R G. and Price, S.D. (1975) AFCRL Infrared Sky Survey, Volume I., Catalog of Observations at 4, 11, and 20 Microns, Environmental Research Papers, No. 522, AFCRL-TR-75-0373.

^{28.} Nadile, R. M., AFGL, Private communications.

6. CONCLUSIONS

A liquid-helium-cooled, long-wavelength infrared (LWIR) spectrometer successfully obtained the first measurements of the infrared spectrum of the upper-atmospheric emissions between 7 and 24 μ m.

All components of the rocket system performed well; however, the apogee obtained was approximately 9 km lower than predicted. This caused the clamshell closing and parachute deployment functions to be initiated at altitudes too low for completely successful execution. As a result, the spectrometer was damaged upon ground impact.

Excellent data were obtained from the LWIR spectrometer which was the primary payload sensor. The 3914 Å photometer provided a good measure of the intensity of the auroral form which was penetrated; however, the dust detector malfunctioned.

During an IBC II aurora, data were obtained on the $15-\mu\,m$ carbon dioxide emission from 65 to 150 km and on the 9.6- $\mu\,m$ ozone emission between 45 and 100 km. The data showed that the most accurate models at the time of the flight were in error by as much as two orders of magnitude leading to a revision of the theoretical predictions. In addition, emission data on the long wavelength wing of the 6.3- $\mu\,m$ water band between 6.7 and 7.6 $\mu\,m$ were obtained between 50 and 75 km.

Above 100 km, significant unidentified emission was observed at 9.3 μ m with weaker features at 6.9, 7.3, 8.0, 11.1, and 12.3 μ m. All these emissions showed a marked asymmetry between ascent and descent. Identification of the species which produce this observed radiation has not been accomplished at this time. Several possible sources have been investigated including atmospheric and exoatmospheric radiators along with contaminants introduced by the measuring system or rocket exhaust. A complete investigation of the spectrometer optics and electronics has shown the emissions are not internally generated and are caused by an external source or sources.

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Appendix A

Spectral Data Scans Obtained During the Flight

This appendix contains all the usable data obtained during the flight. Individual scans are shown at altitudes below 104 km in Figures A1 through A11 (96.3 to 103.2 km) during ascent and in Figures A77 through A151 (102.9 to 45.6 km) during descent. At the higher altitudes, scans averaged over 2.5 km increments are shown in Figures A17 through A43 during ascent and Figures A44 through A76 during descent.

For individual scans, the scan number, time, and altitude are shown at the top of the figures. The inclusive scan numbers, times, and altitudes are shown for the averaged scans. The numbers in parentheses indicate the number of usable scans and the total number of scans in the 2.5-km increment being averaged. The spectral intensity values are in W cm $^{-2}$ sr $^{-1}$ μm $^{-1}$, with the exponent shown in the upper right-hand corner of the figure. The clipping of the 15- μ m CO $_2$ below 67 km is caused by amplifier saturation.

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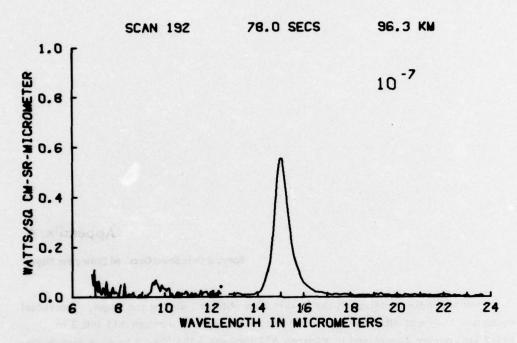


Figure A1

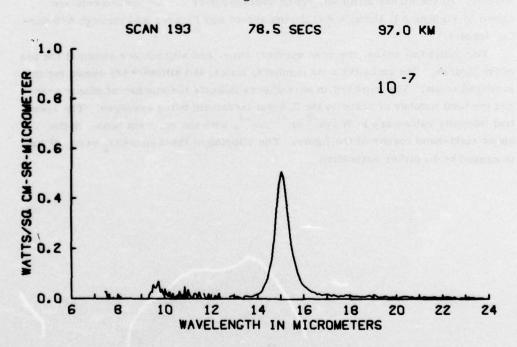


Figure A2

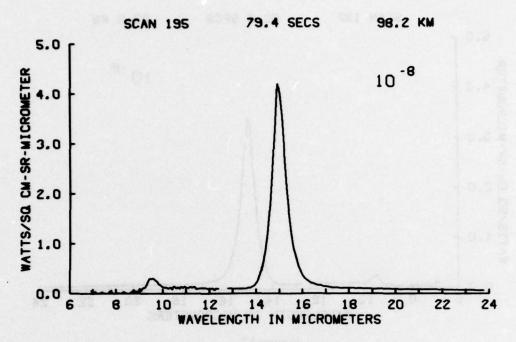


Figure A3

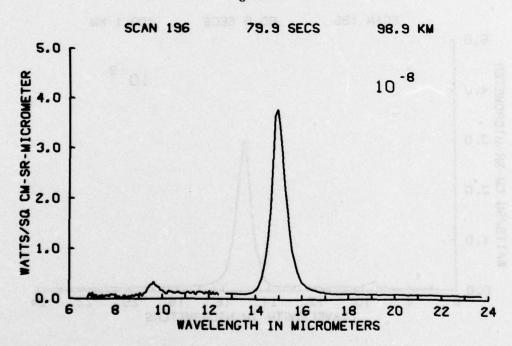


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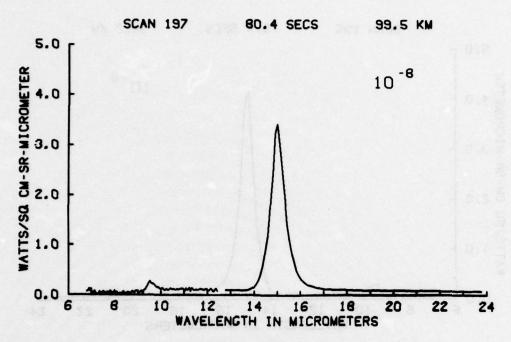


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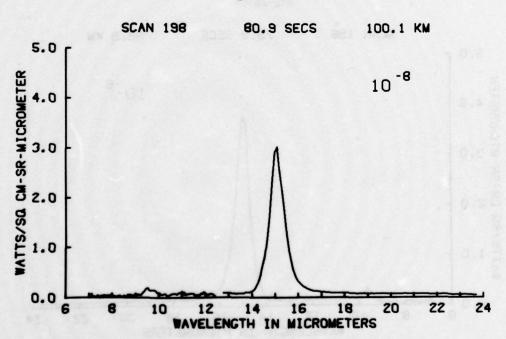


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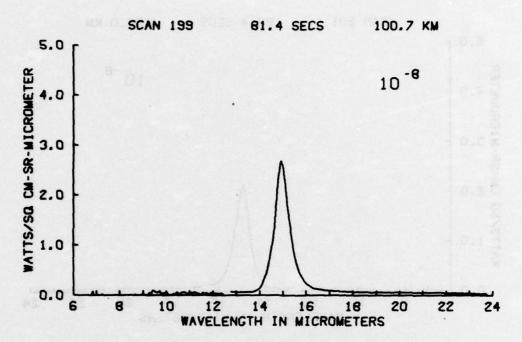


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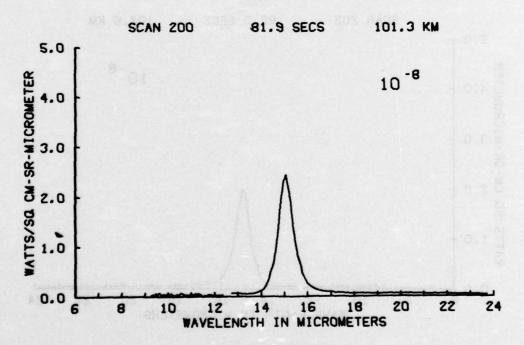


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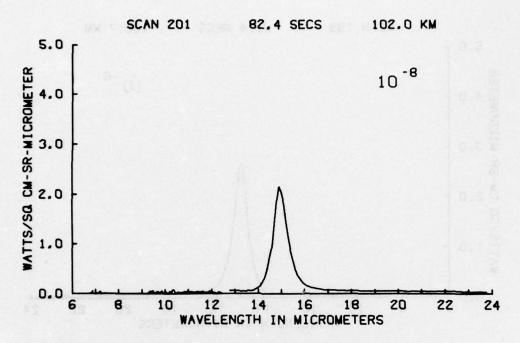


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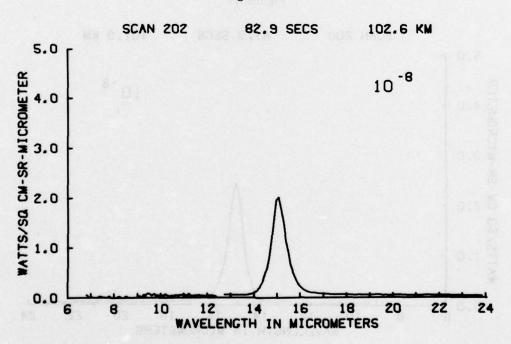


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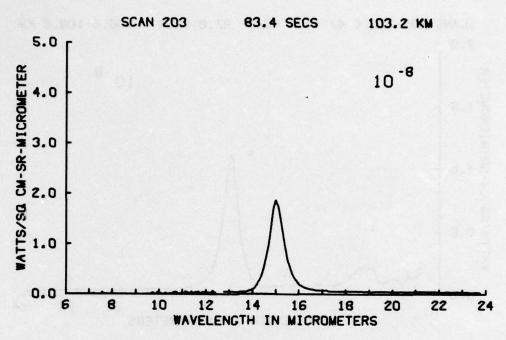


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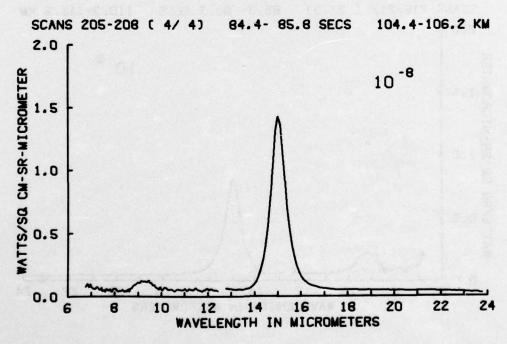


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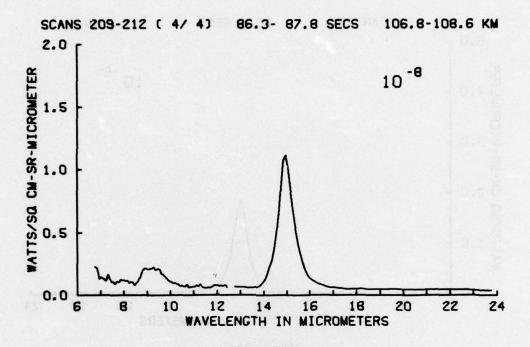


Figure A13

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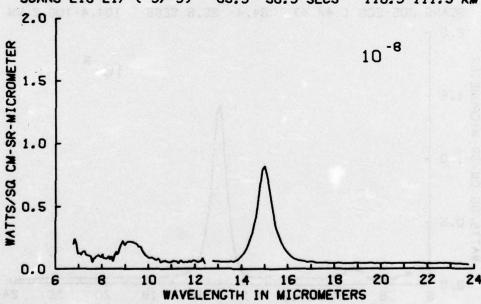


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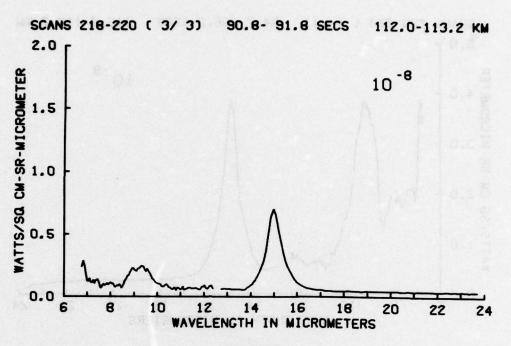


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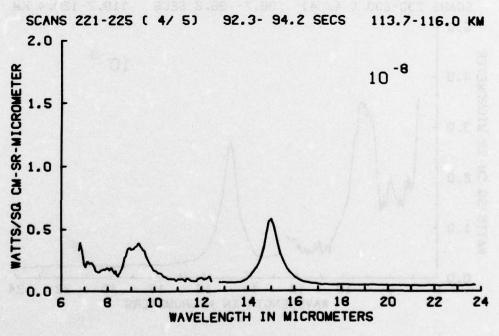


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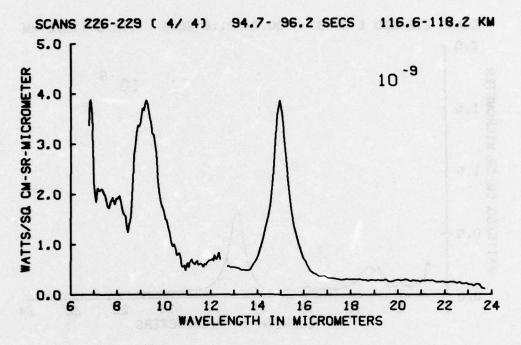


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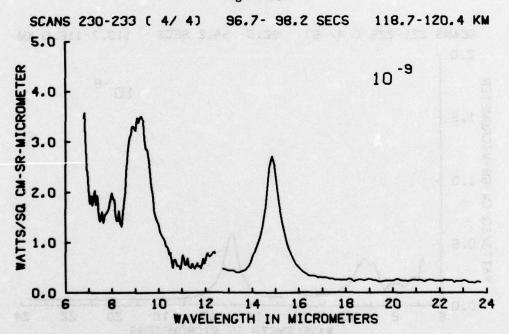


Figure A18

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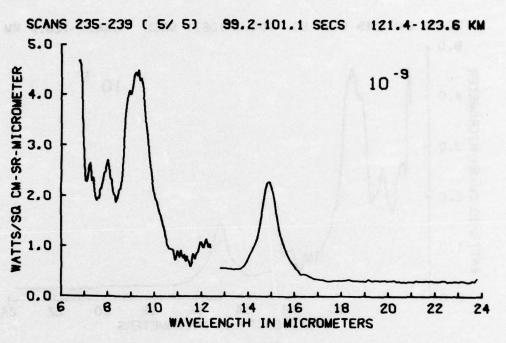


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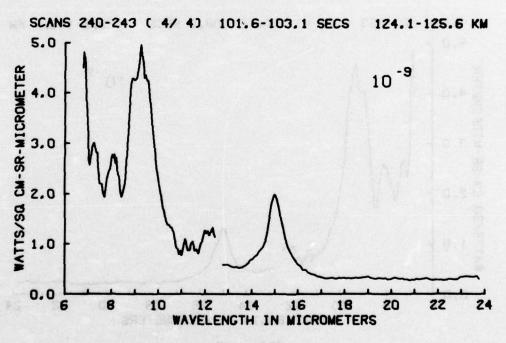


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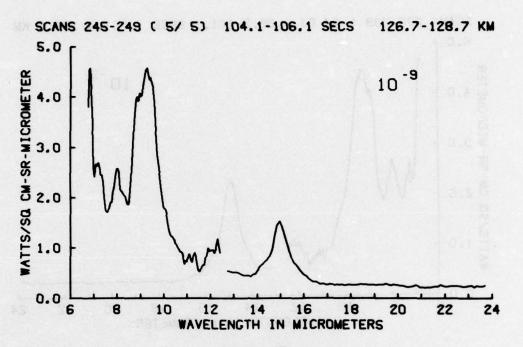


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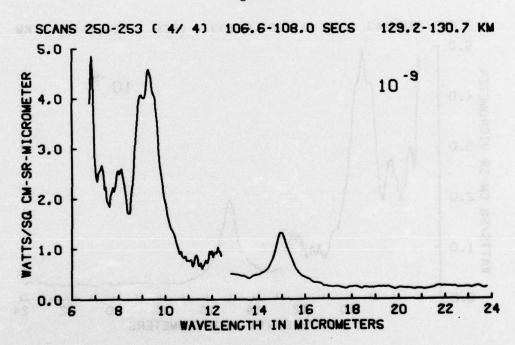


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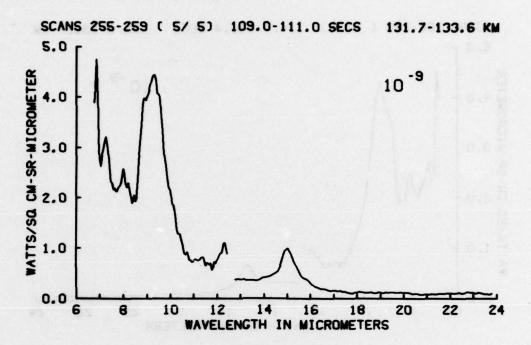


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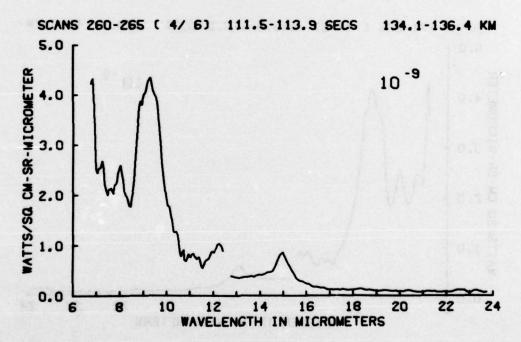


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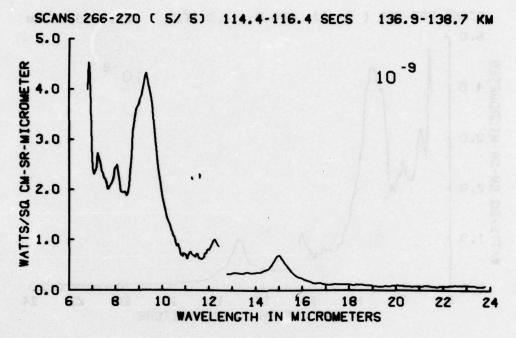


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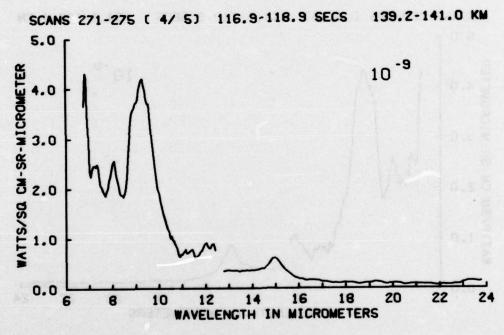


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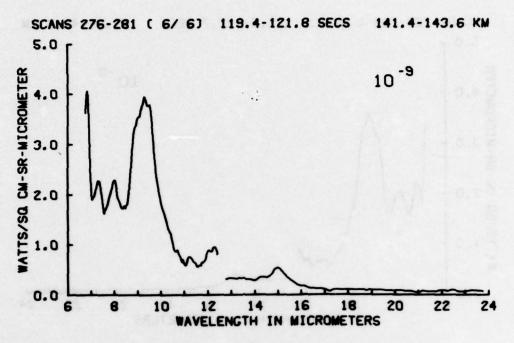


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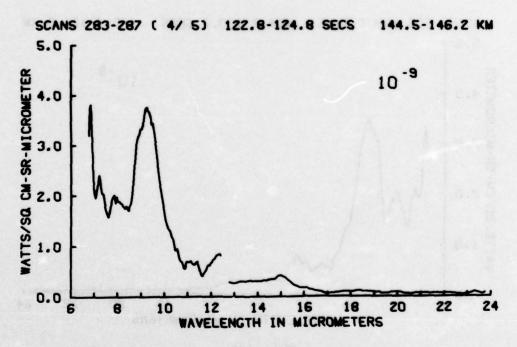


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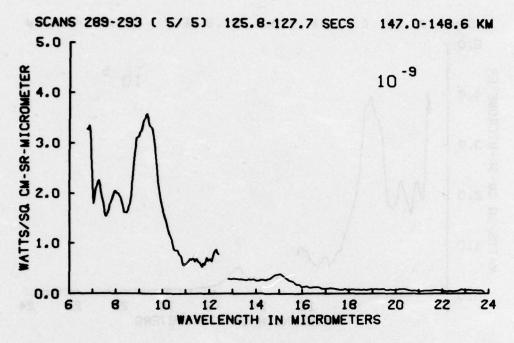


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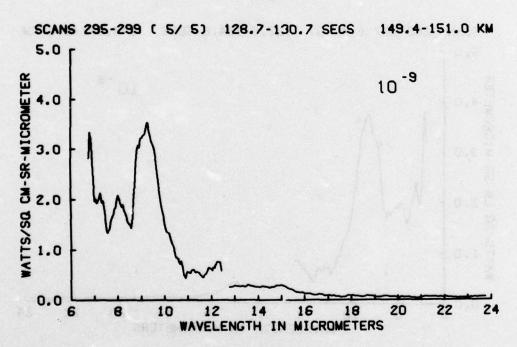


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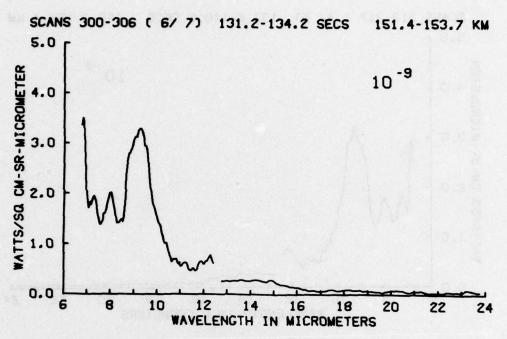


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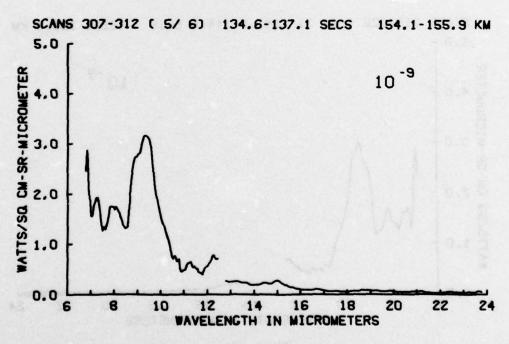


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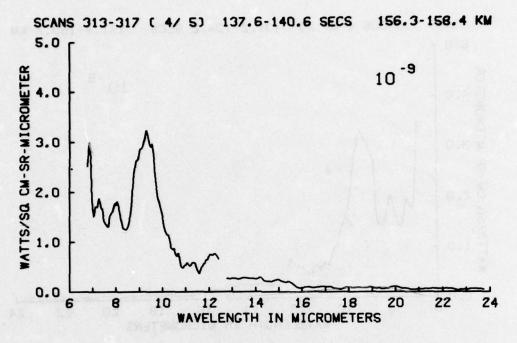


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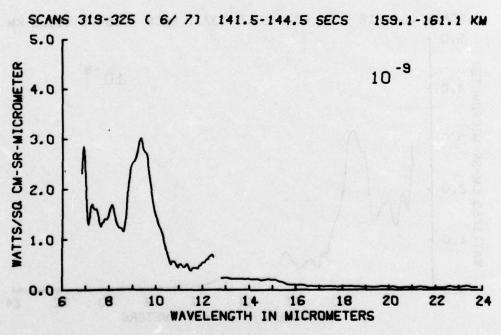


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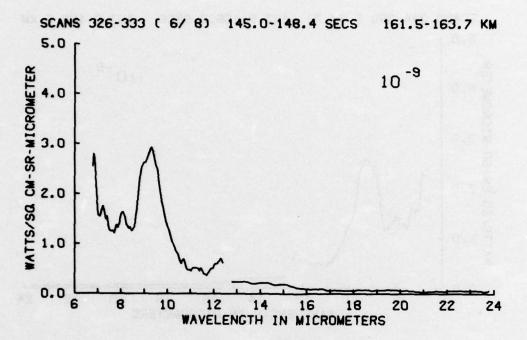


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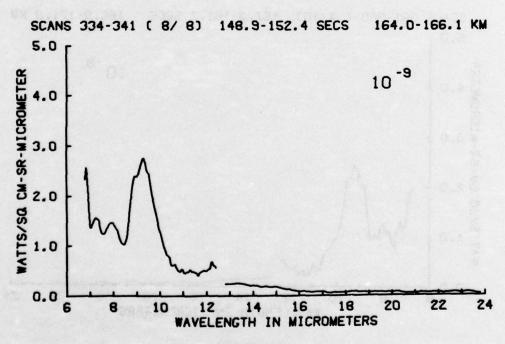


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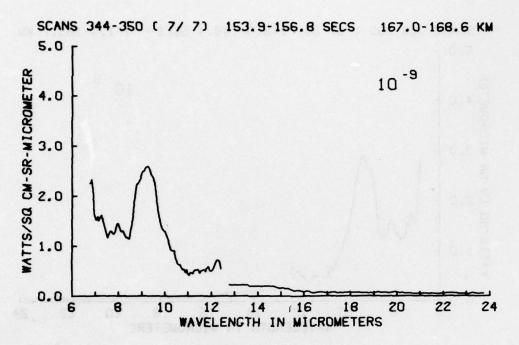


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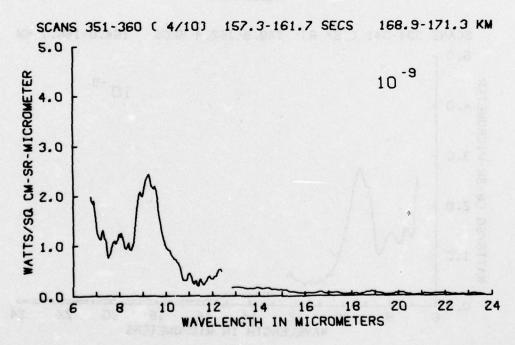


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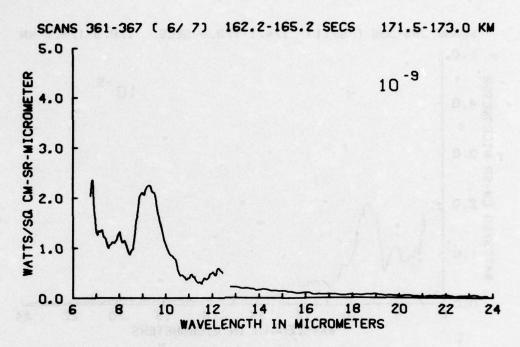


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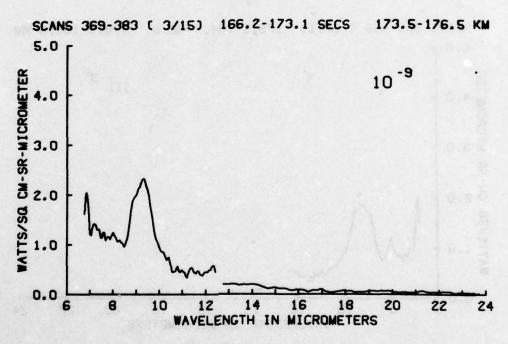


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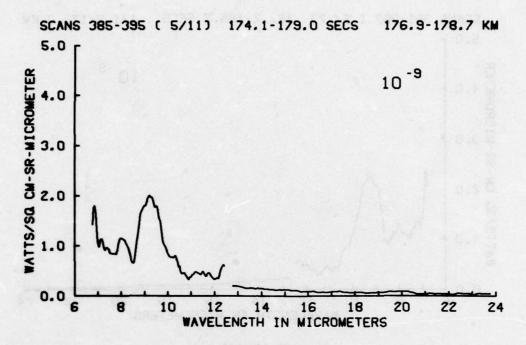


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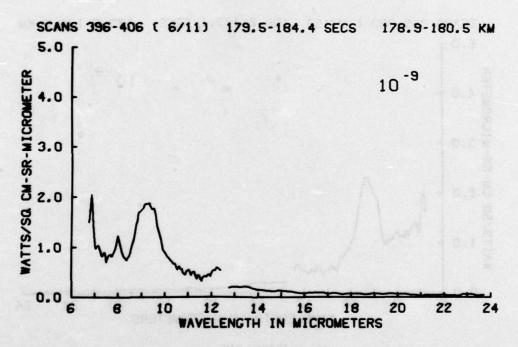


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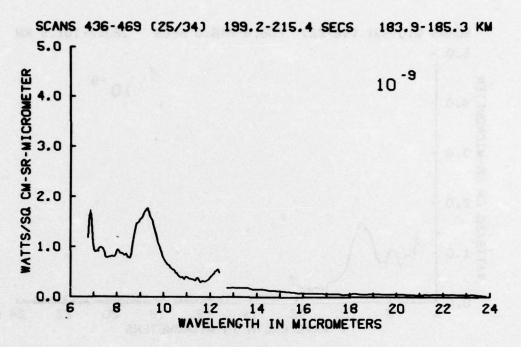


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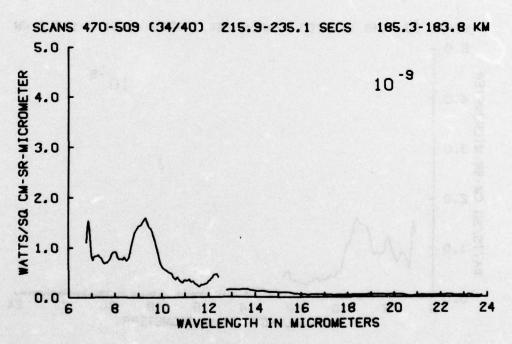


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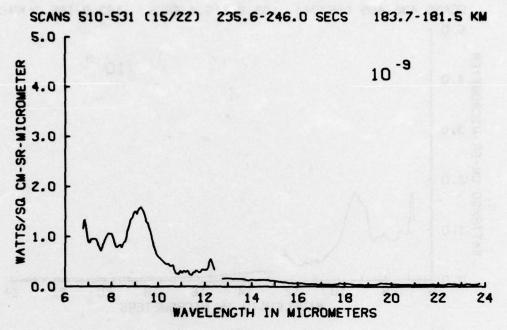


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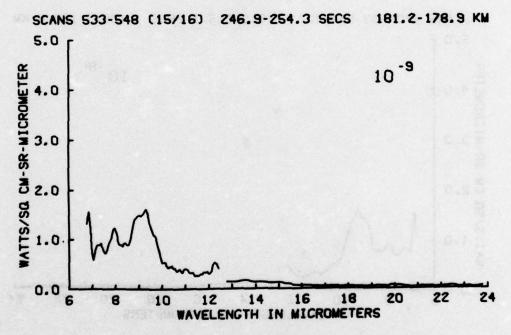


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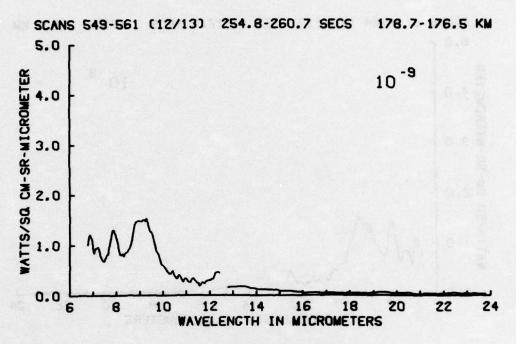


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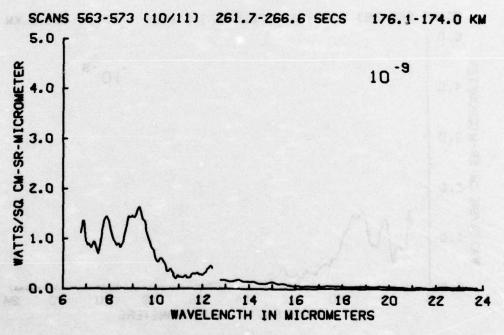


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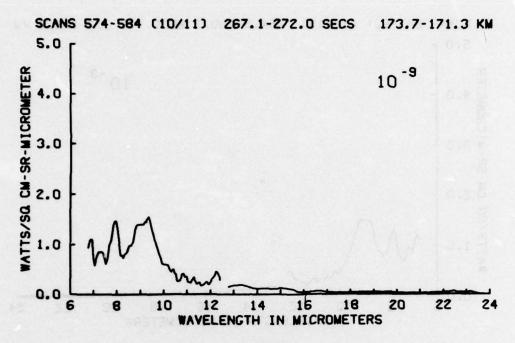


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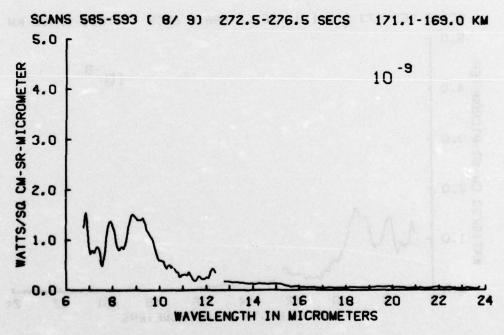


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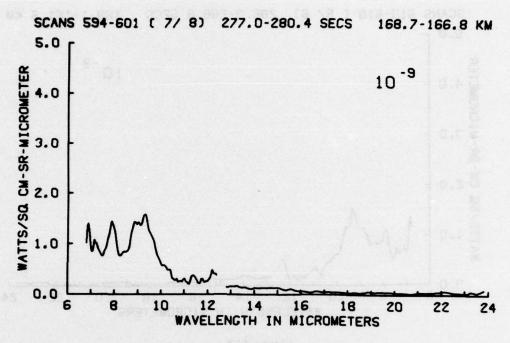


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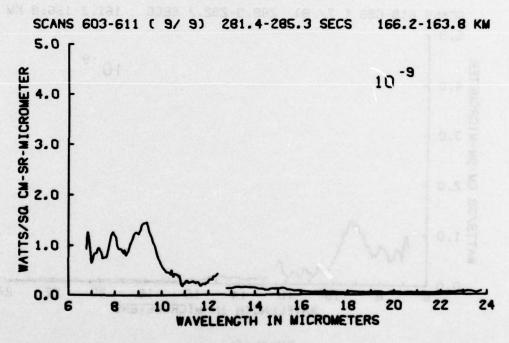


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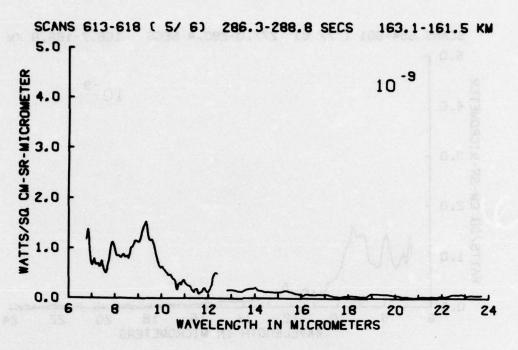


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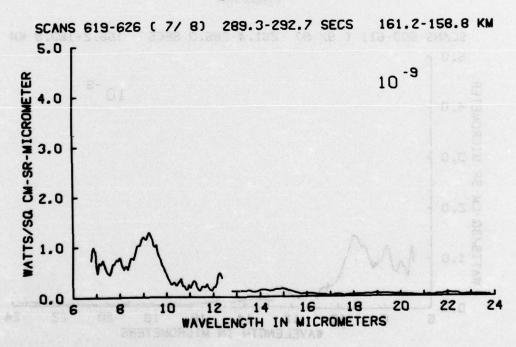


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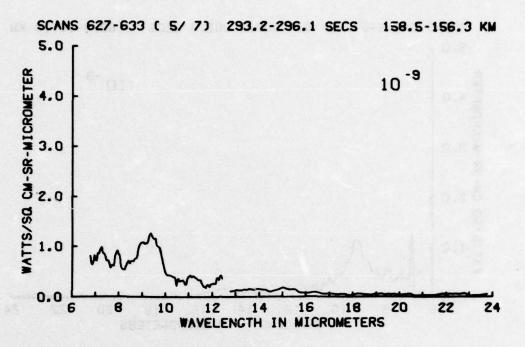


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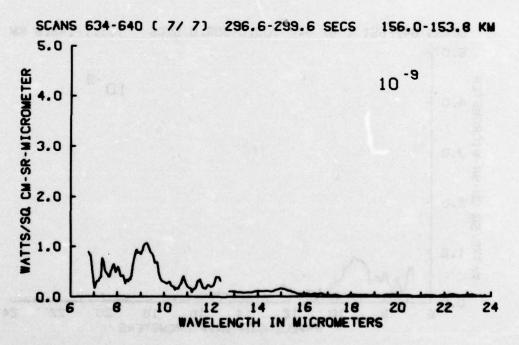


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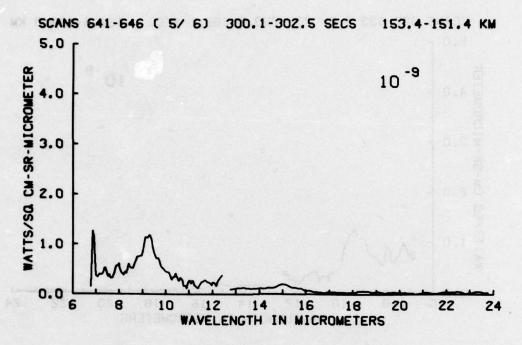
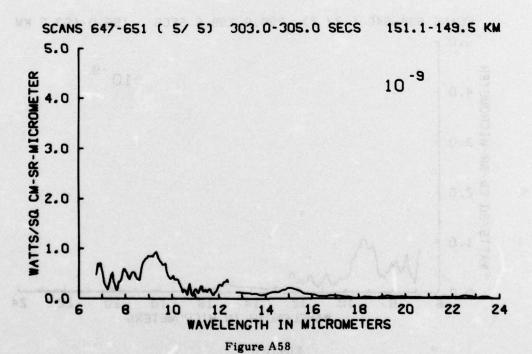


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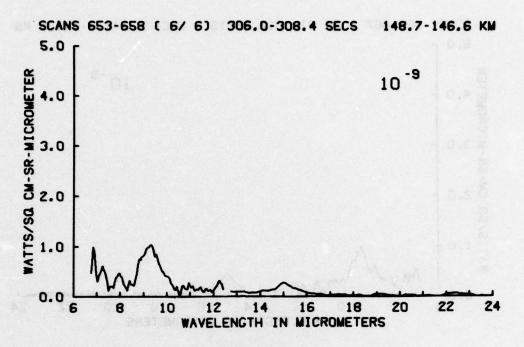


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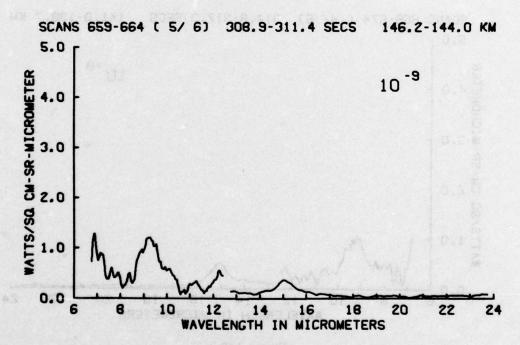


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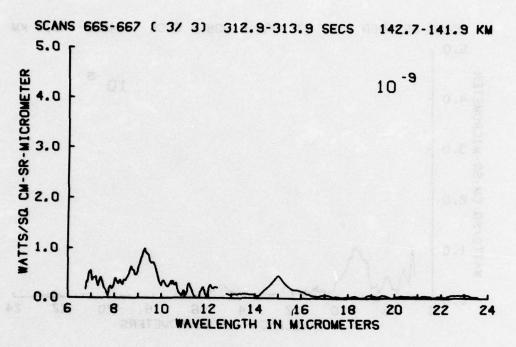


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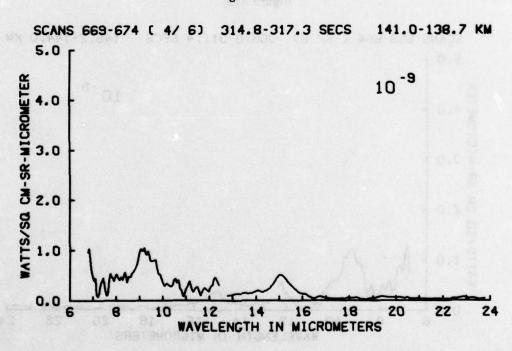


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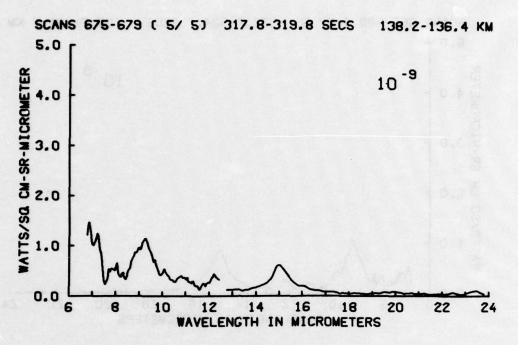


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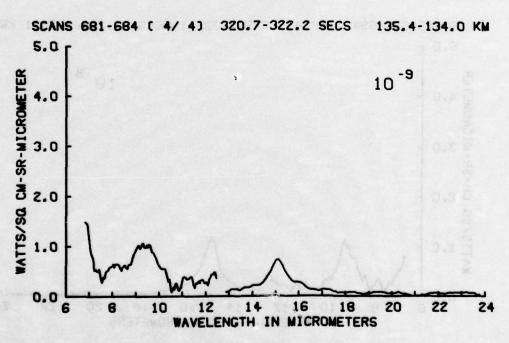


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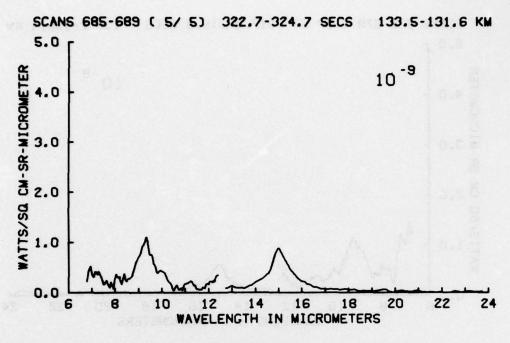


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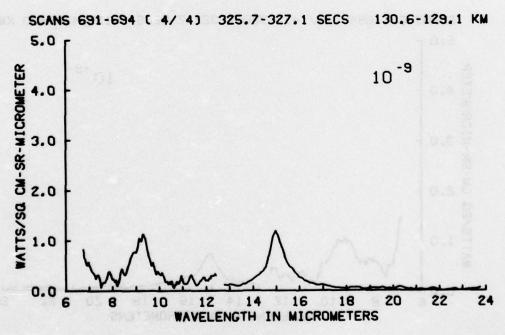


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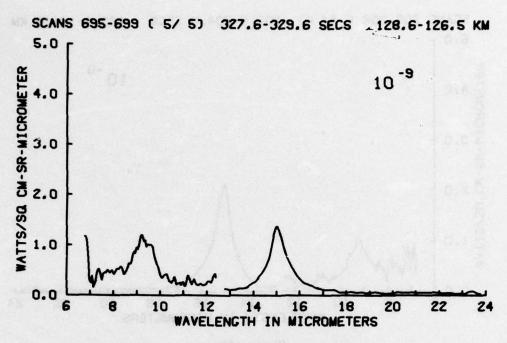


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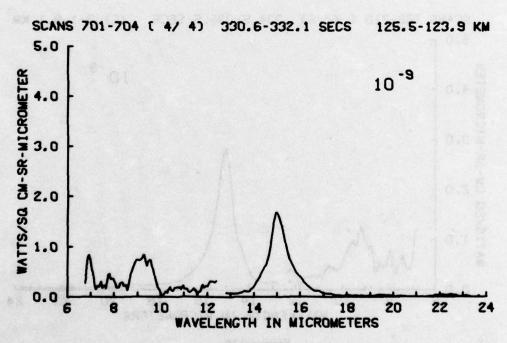


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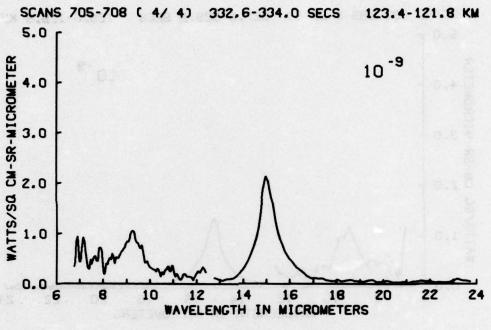


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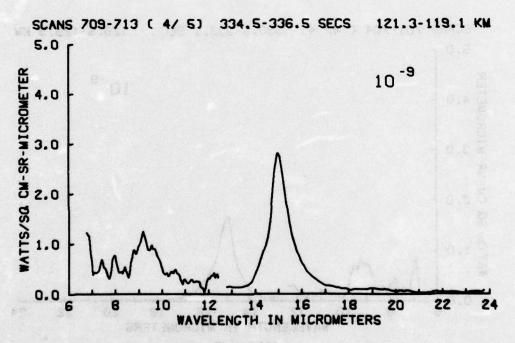


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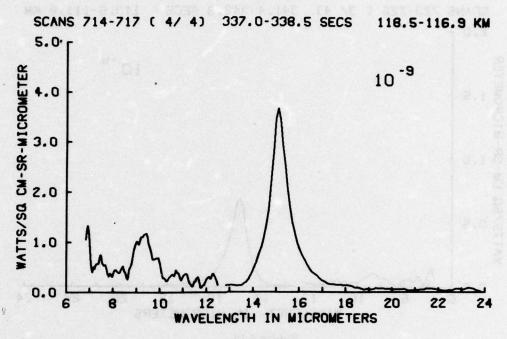


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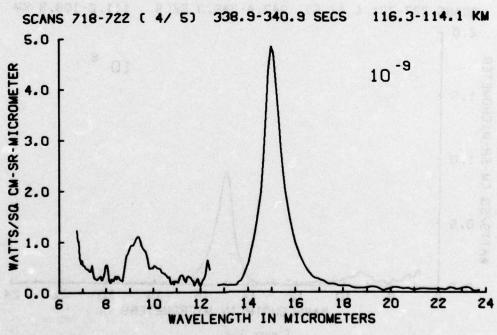


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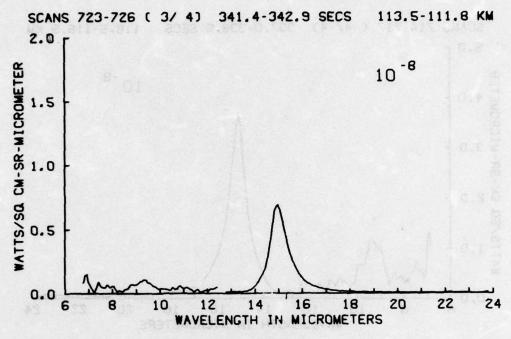


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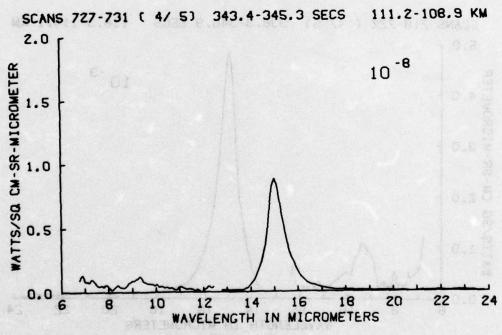


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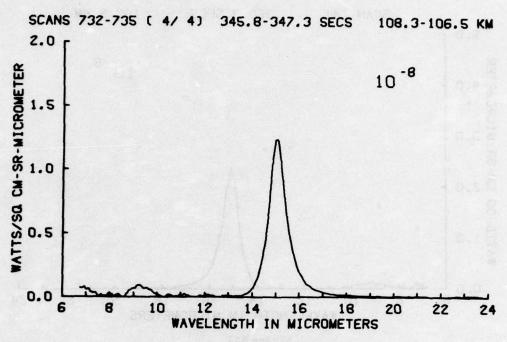


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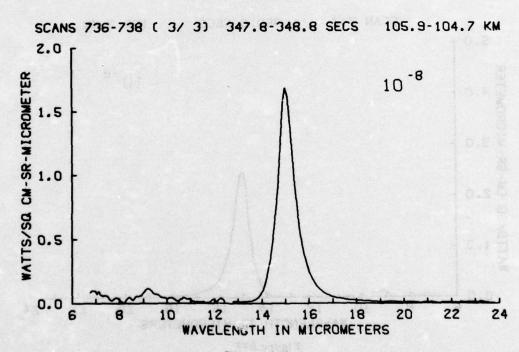


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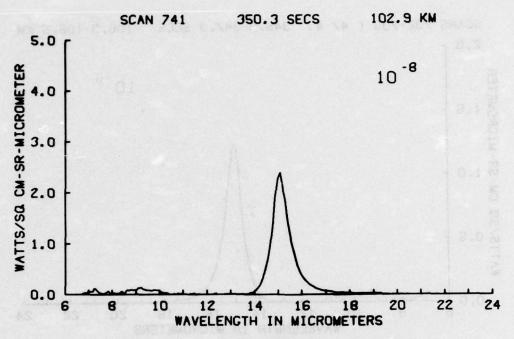
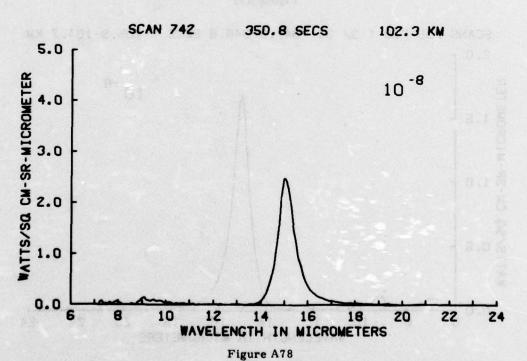
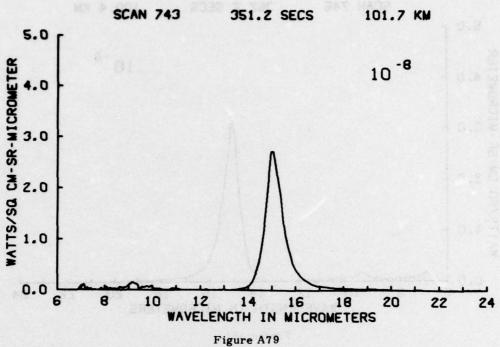


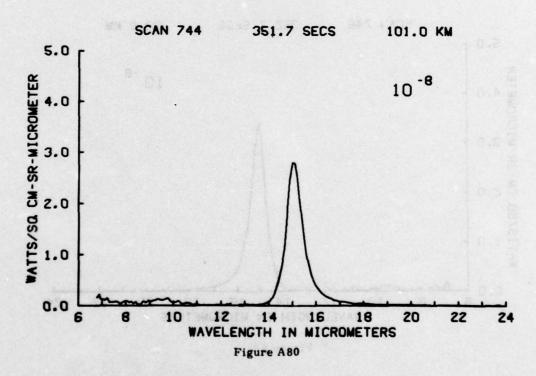
Figure A77



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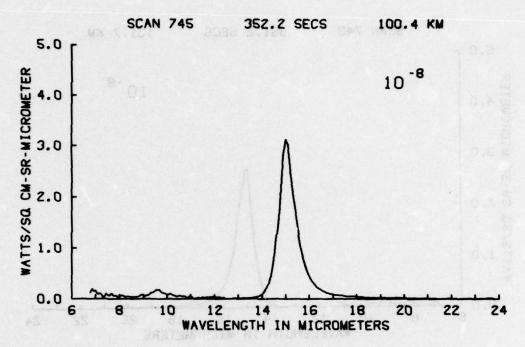


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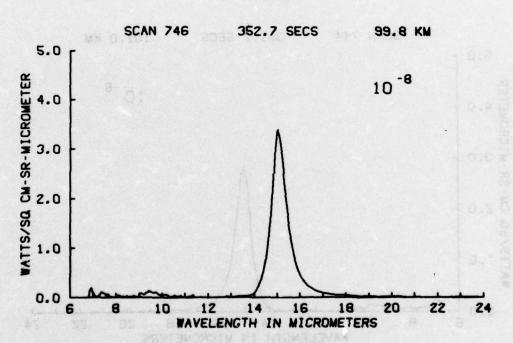


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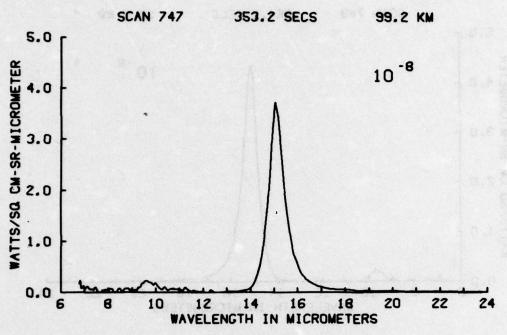


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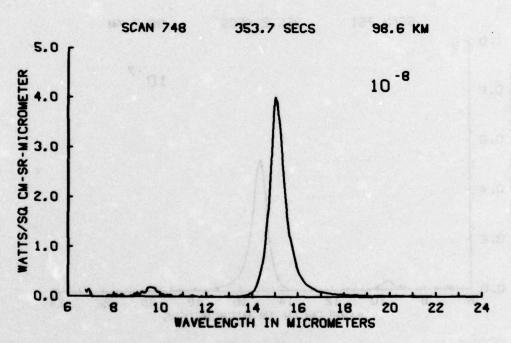


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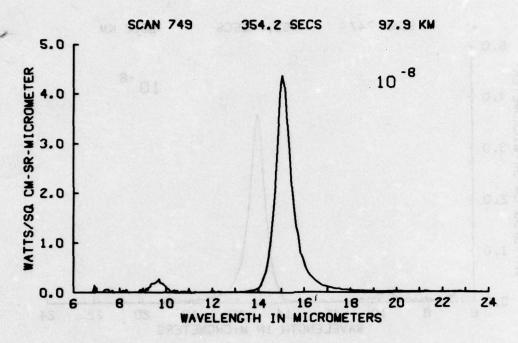


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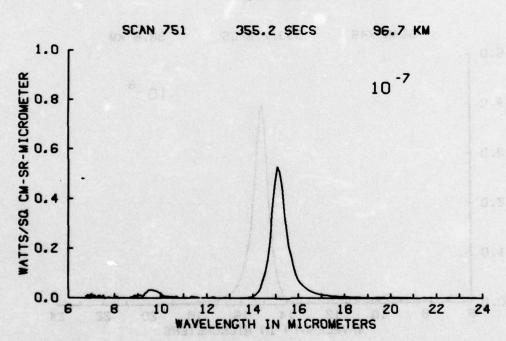


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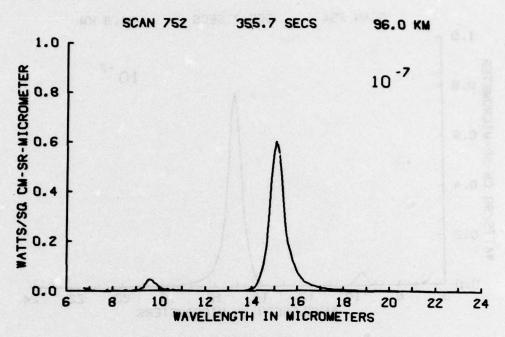


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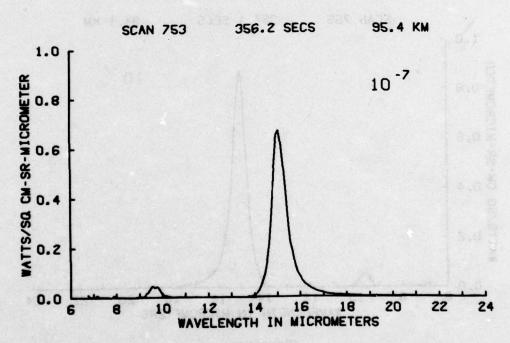


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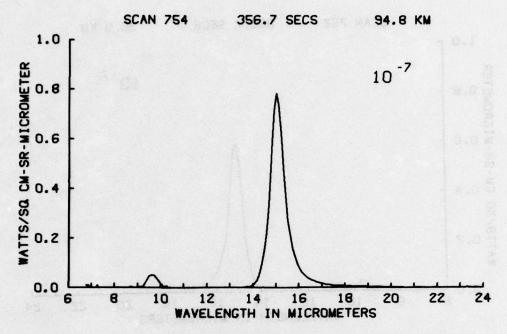


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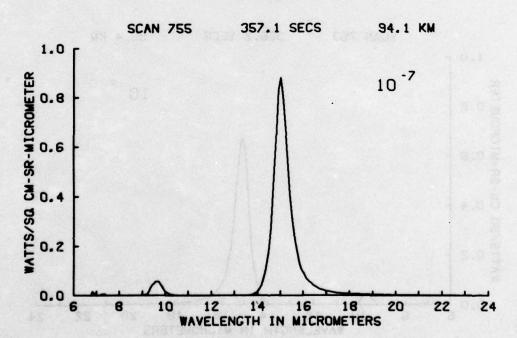


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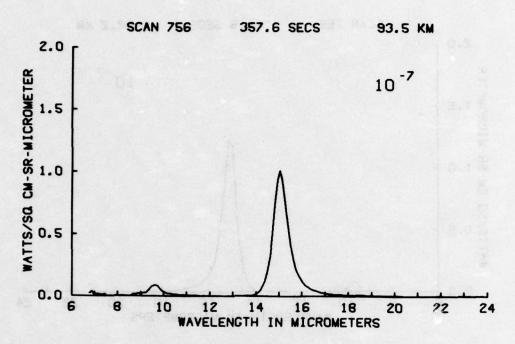


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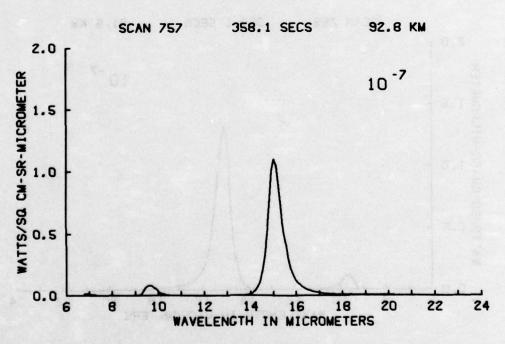


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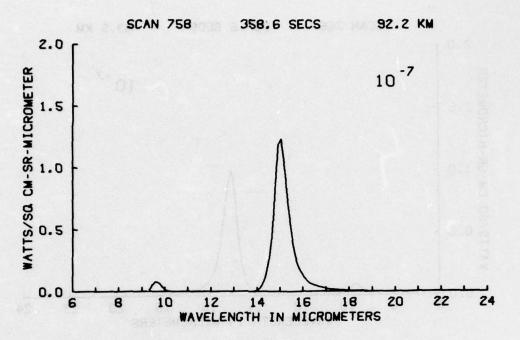


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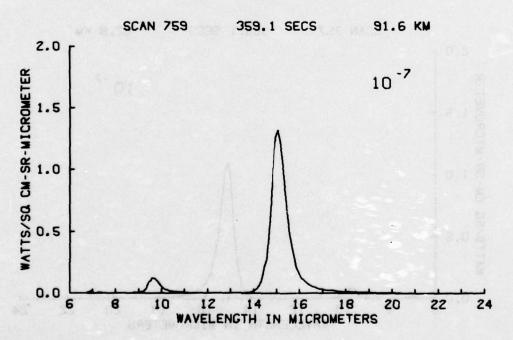


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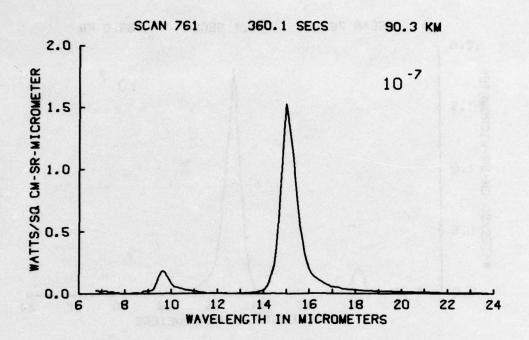


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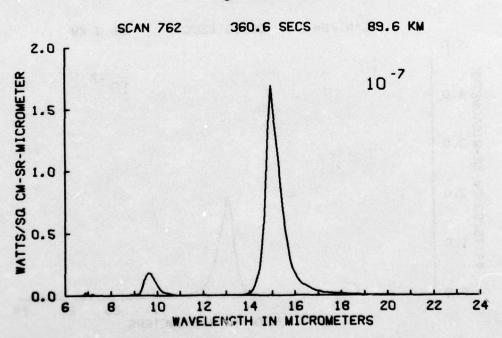


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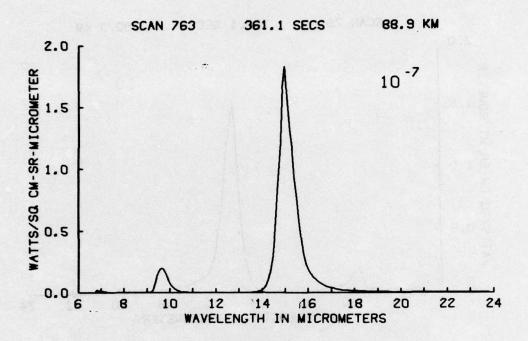


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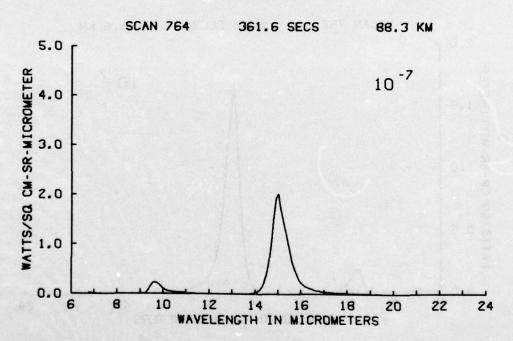


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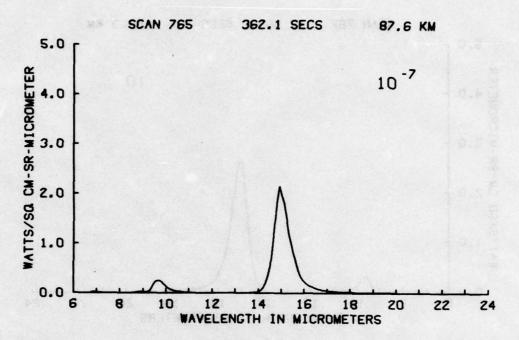


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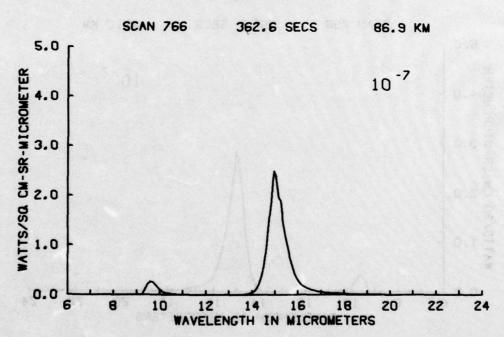


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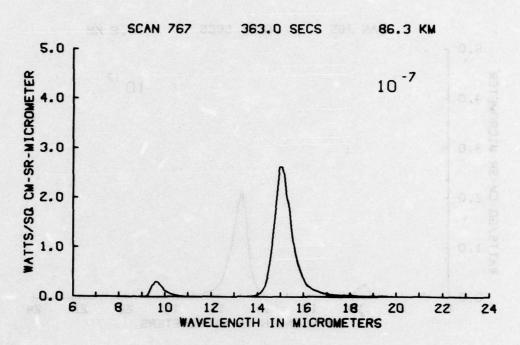


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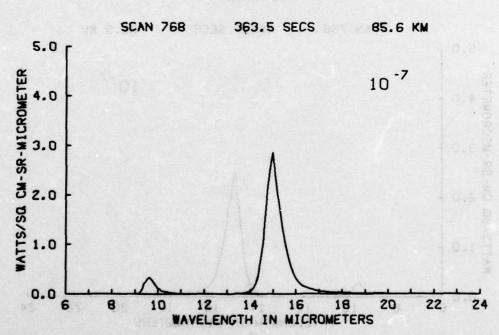


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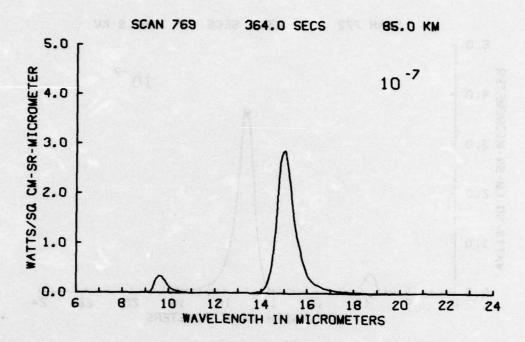


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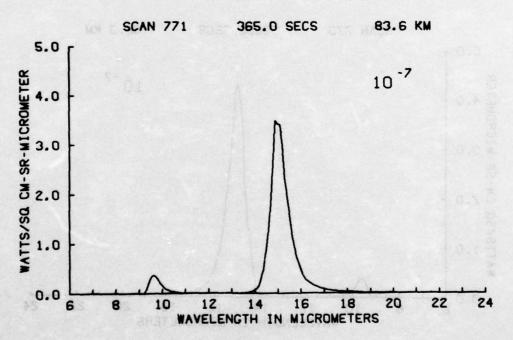


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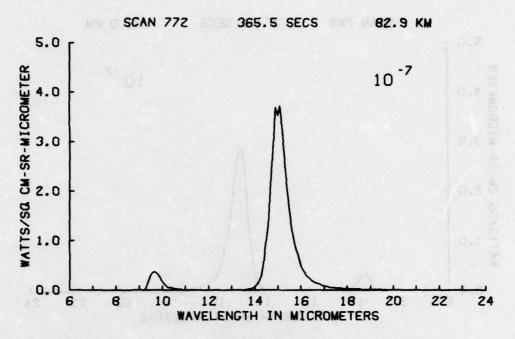


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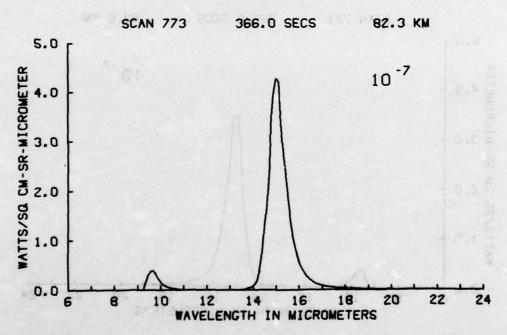


Figure A106

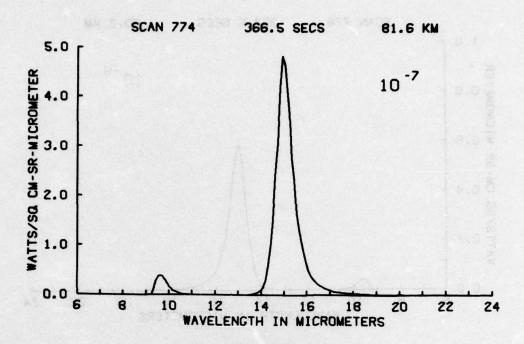


Figure A107

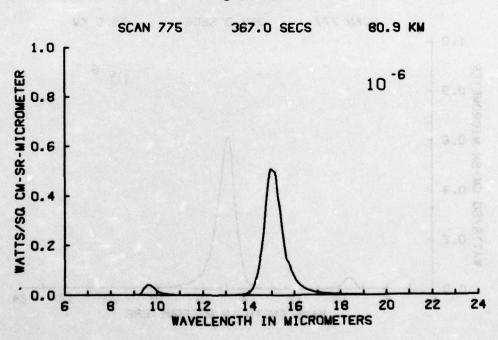


Figure A108

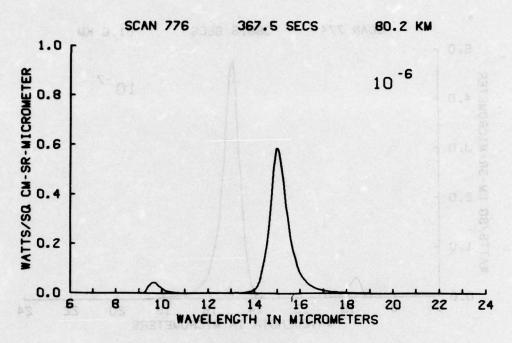


Figure A109

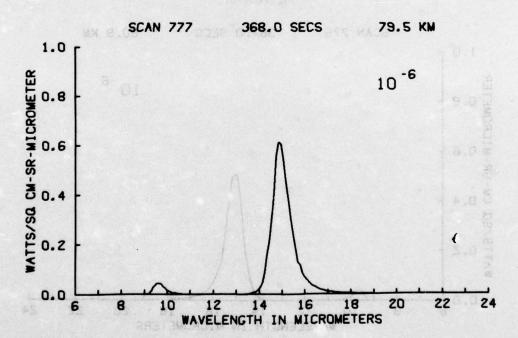


Figure A110

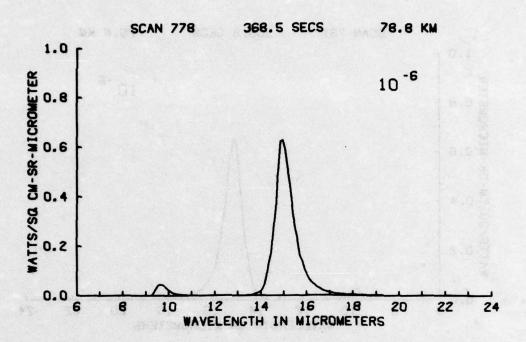


Figure A111

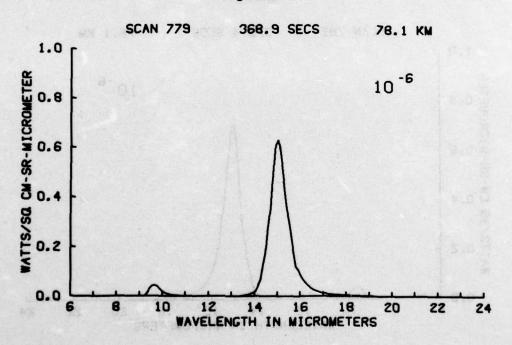


Figure A112

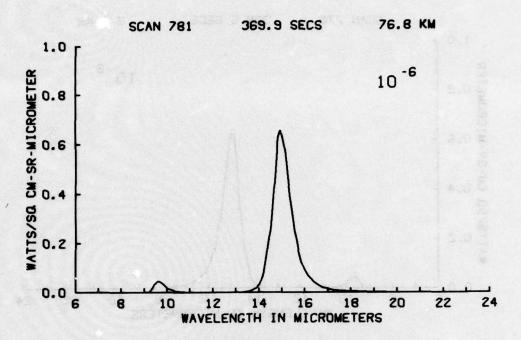


Figure A113

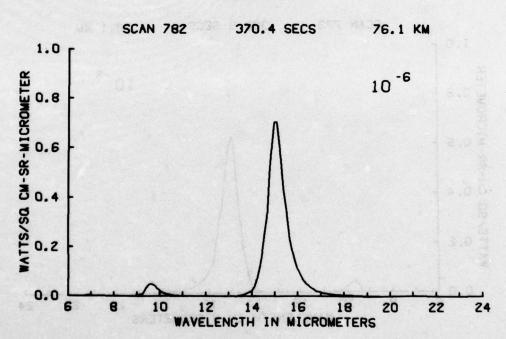


Figure A114

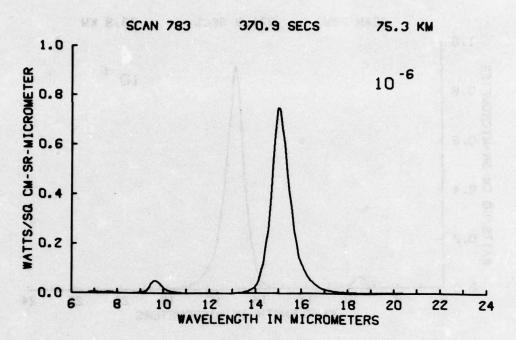


Figure A115

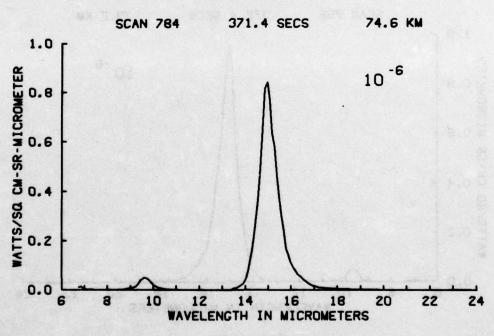


Figure A116

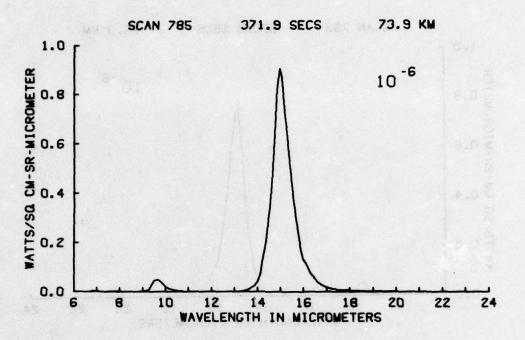


Figure A117

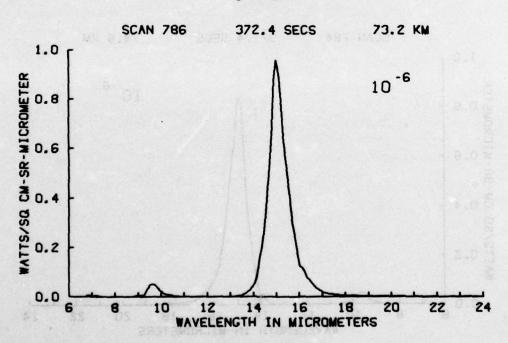


Figure A118

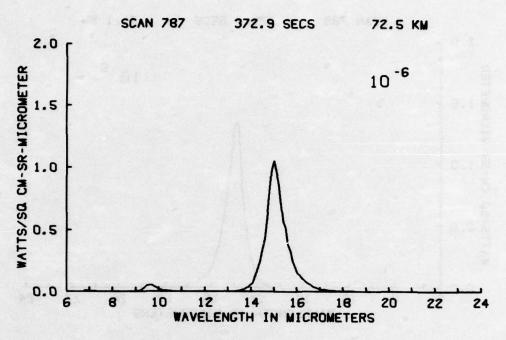


Figure A119

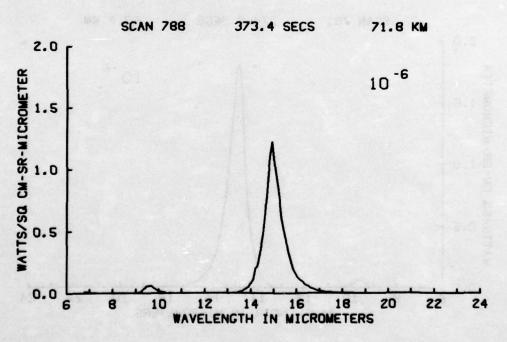


Figure A120

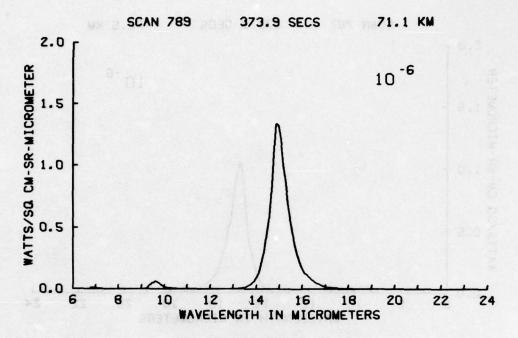


Figure A121

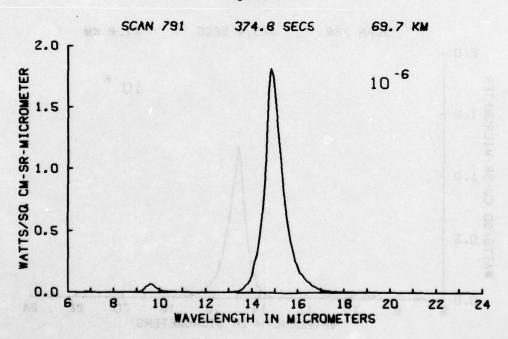
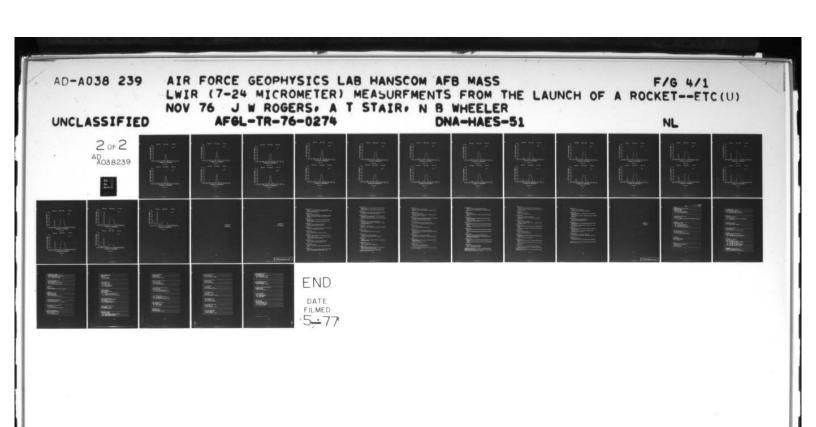


Figure A122



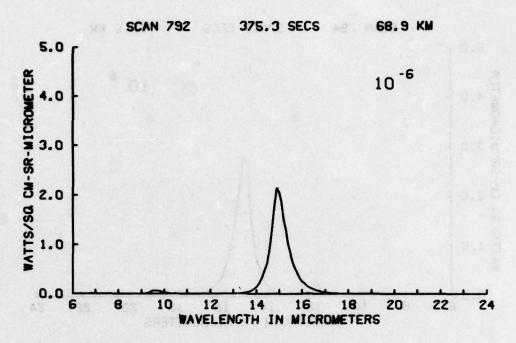


Figure A123

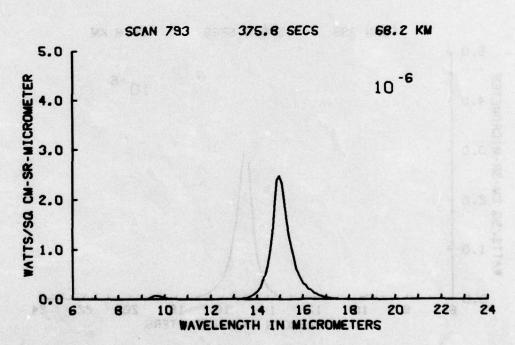


Figure A124

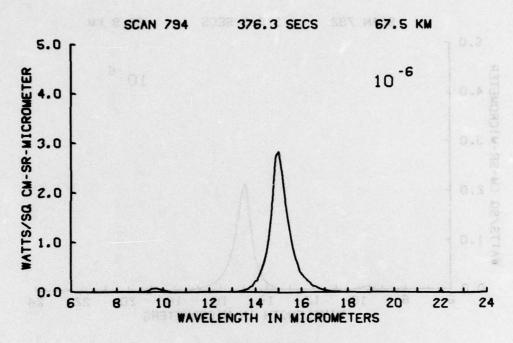


Figure A125

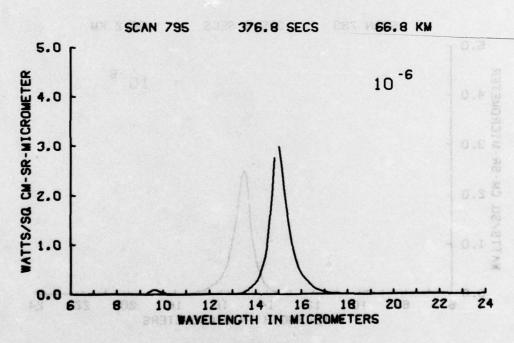


Figure A126

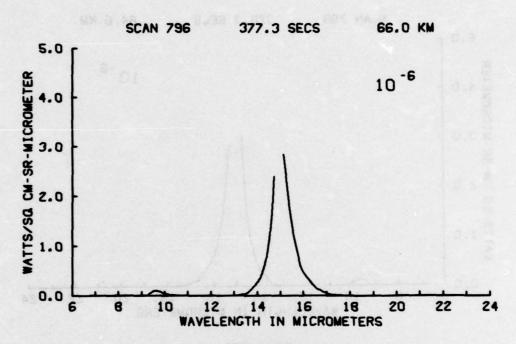


Figure A127

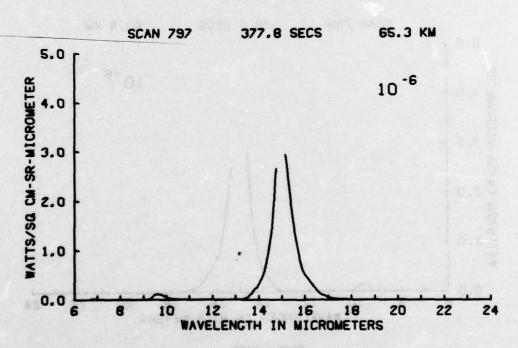


Figure A128

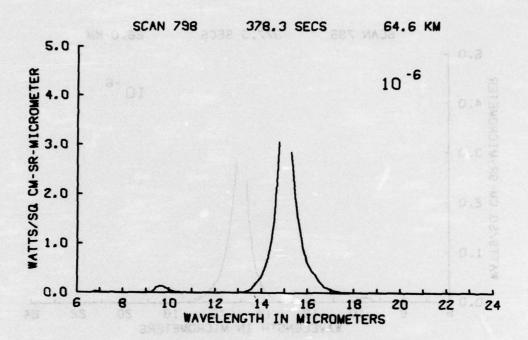


Figure A129

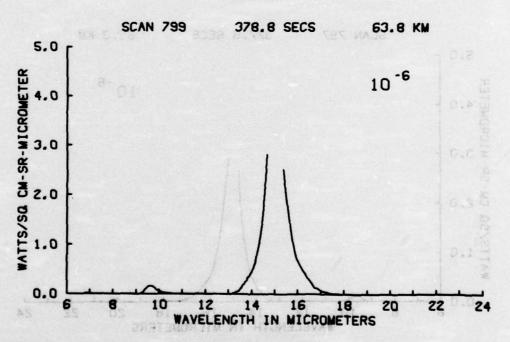


Figure A130

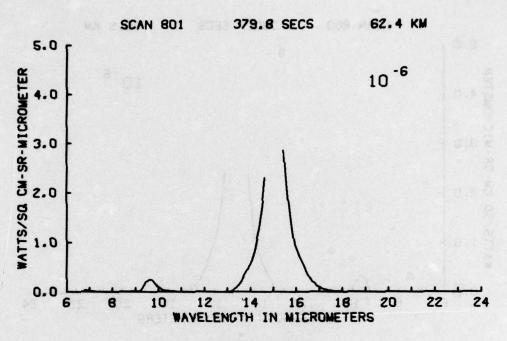


Figure A131

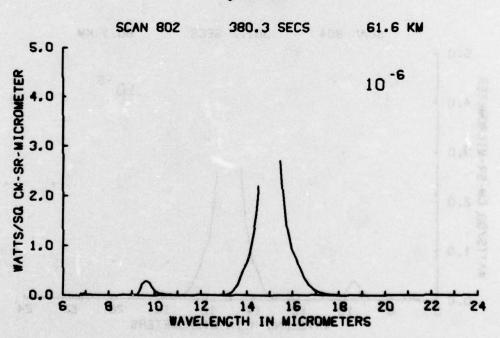


Figure A132

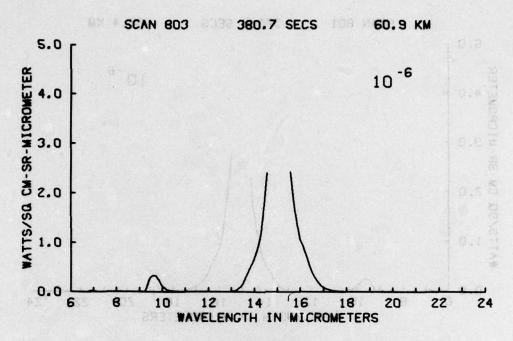


Figure A133

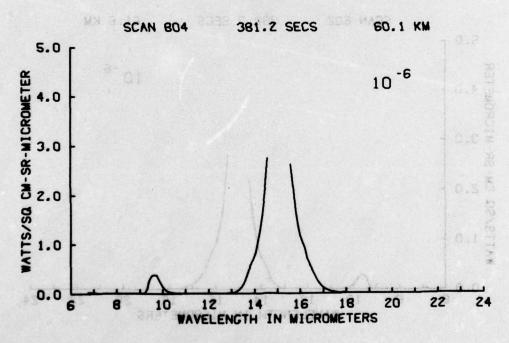


Figure A134

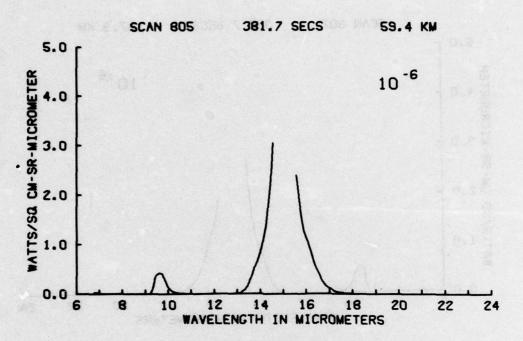


Figure A135

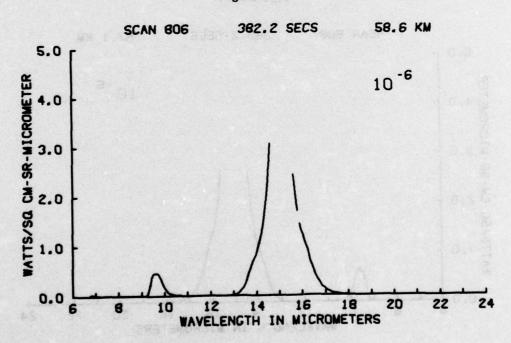


Figure A136

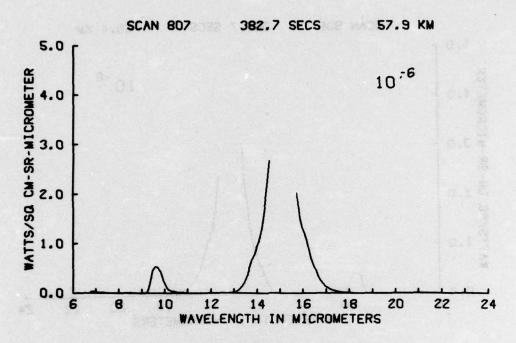


Figure A137

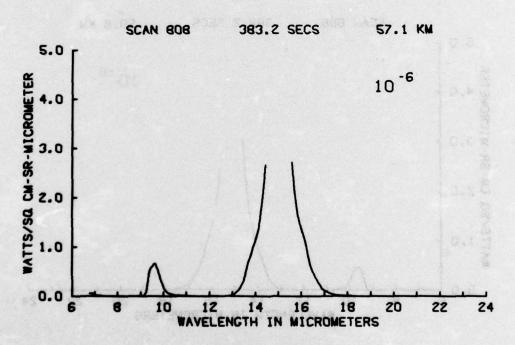


Figure A138

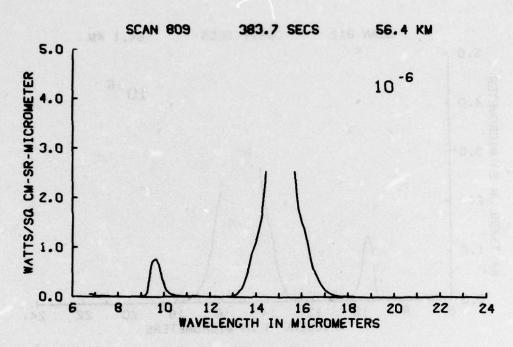


Figure A139

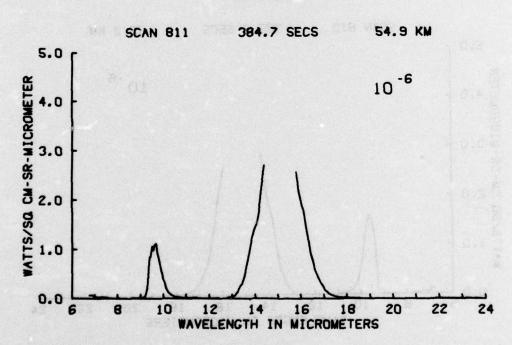


Figure A140

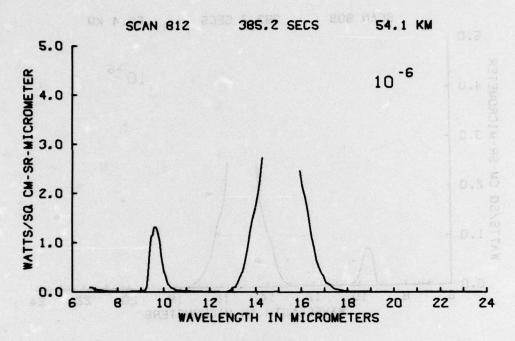


Figure A141

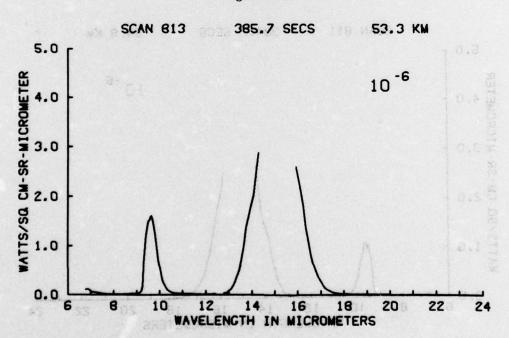


Figure A142

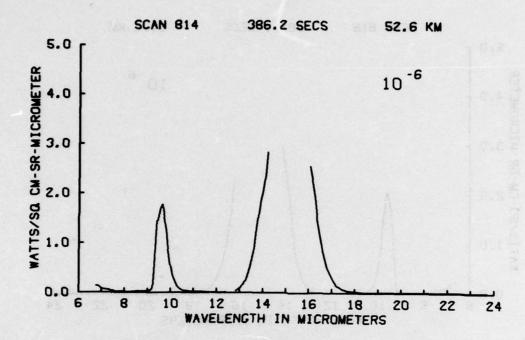


Figure A143

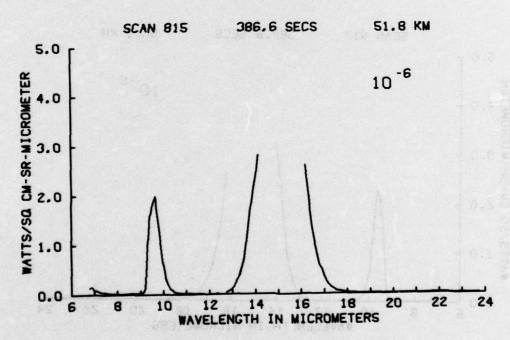


Figure A144

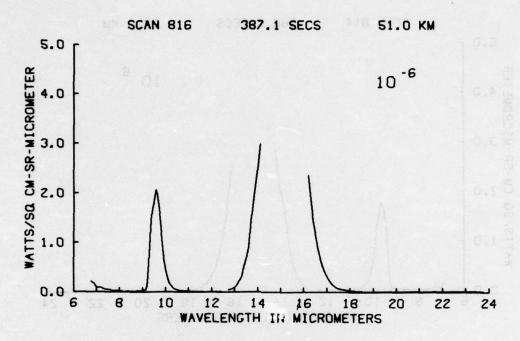


Figure A145

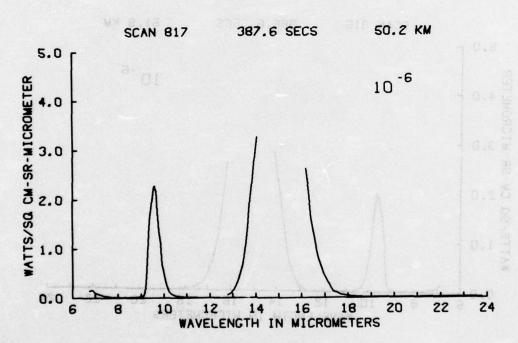


Figure A146

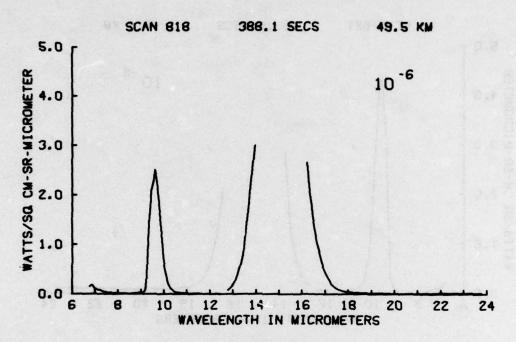


Figure A147

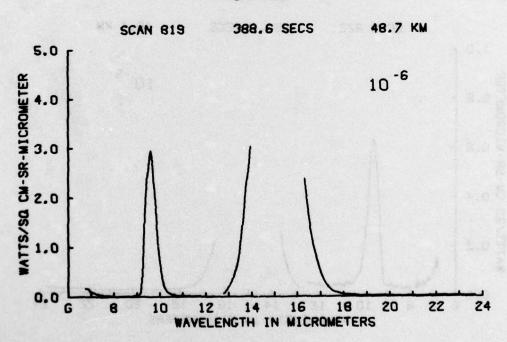


Figure A148

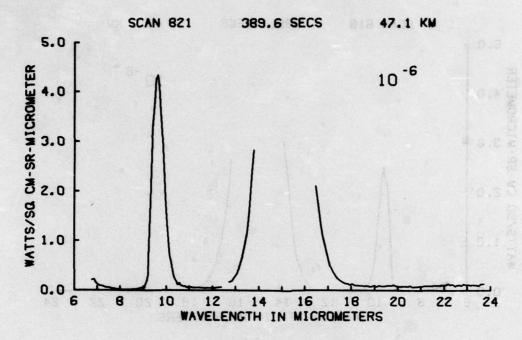


Figure A149

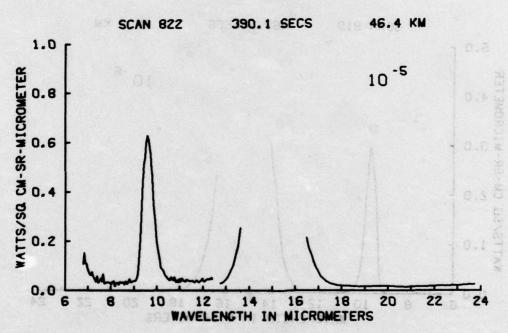


Figure A150

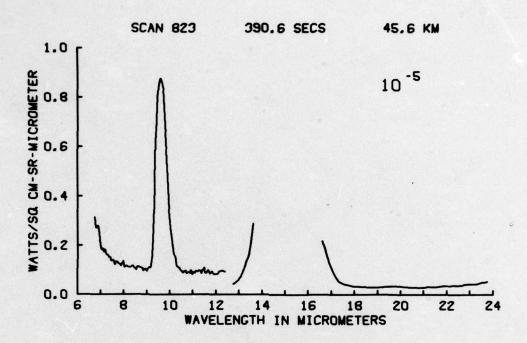


Figure A151

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Appendix B
List of HAES Reports

Rocket Launch of a SWIR Spectrometer into an Aurora (ICECAP 72)

D.J. Baker, C.L. Wyatt, W.R. Pendelton, Jr. — Utah State, J.C. Ulwick — AFCRL;

AFCRL-TR-74-0077, February 1974.

HAES Report No. 2

Analysis of HAES Results: ICECAP 72

W.P. Reidy, T.C. Degges, O.P. Manley, H.J. Smith, J.W. Carpenter — Visidyne, A.T. Stair, J.C. Ulwick — AFCRL, D.J. Baker — Utah State; DNA 3247F, April 1974.

HAES Report No. 3

Rocket Instrumentation for ICECAP 73A, Auroral Measurements Program — Black Brant 18.205-1

D.A. Burt, C.S. Davis - Utah State; AFCRL-TR-74-0195, February 1974.

HAES Report No. 4

Data Reduction and Auroral Characterization for ICECAP

1. Kofsky — Photometrics; DNA 3511F, January 1975.

HAES Report No. 5

ICECAP Analysis: Energy Deposition and Transport in the Auroral Ionosphere R.D. Sears — Lockheed; DNA 3566F, November 1974

HAES Report No. 6

Auroral Simulation Studies

D. Archer - Mission Research Corp; DNA 3567T, April 1975.

HAES Report No. 7

ICECAP 72 — Rocket Measurement Program for Investigation of Auroral IR Emissions — Black Brant 17.110-3.

D.A. Burt, G.D. Allred, J.C. Kemp, L.C. Howlett, E.F. Pound, G.K. LeBaron – Utah State; AFCRL-TR-75-0001, September 1974.

HAES Report No. 8

Design and Calibration of a Rocket-Borne Electron Spectrometer P.C. Neal — Utah State; AFCRL-TR-74-0629, December 1974.

HAES Report No. 9

Near Infrared Auroral Spectra

D. Baker, A. Steed, R. Huppi - Utah State; AFCRL-TR-75-0010, December 1974.

HAES Report No. 10

Arctic Code Electron Deposition Theory with Application to Project EXCEDE P.W. Tarr -- Mission Research Corp; MRC-R-173, February 1975.

Rocketborne Instrumentation for the Measurement of Electric Fields — Paiute Tomahawk 10.312-3

L. Carl Howlett, R.J. Bell - Utah State; AFCRL-TR-75-0023, January 1975.

HAES Report No. 12

Auroral Chemistry and Energy Deposition Calculations with the OPTAUR Code H.J.P. Smith, O.P. Manley — Visidyne; VI 266, February 1975.

HAES REPORT No. 13*

Theoretical Evaluation of Vertically Viewing and Earth-Limb Scanning Modes for Rocketborne Earth-Limb Measurements

C.L. Wyatt, R.Y. Han, D.J. Baker — Utah State; AFCRL-TR-75-0072, September 1974.

HAES Report No. 14

Ionospheric Effects Induced by Precipitating Auroral Electrons

J.D. Cladis, G.T. Davidson, W.E. Francis, L.L. Newkirk, M. Walt — Lockheed; LMSC/D454890, February 1975.

HAES Report No. 15

ICECAP 73A - Chatanika Radar Results

M.J. Baron, N.J. Chang -- Stanford Research Institute; DNA 3531T, September 1974.

HAES Report No. 16

Numerical Modeling of Aurora

T. Coffey - Naval Research Labs; NRL-MR-3120, October 1975.

HAES Report No. 17

TMA Payload Field Services and Data Reduction

J.F. Bedinger -- GCA Corp; AFCRL-TR-75-0009, February 1975.

HAES Report No. 18

Auroral NO

E. Hyman, D.J. Strickland, P.S. Julienne, D.F. Strobel — Naval Research Labs; NRL-MR-3070, July 1975.

HAES Report No. 19

Analysis of 4.3µ ICECAF Data

J. Kumer - Lockheed; AFCRL-TR-74-0334, July 1974.

HAES Report No. 20

Cylindrical Langmuir Probe Measurements from Rocket Flights Covering the Period 31 January 1961 through 3 April 1974

R.C. Wilson - Boston College; AFCRL-TR-75-0265, April 1975.

*HAES number omitted from this report. If you are a recipient of this report, please add the appropriate HAES report number to the front cover.

Data System Developed for Project ICECAP 74

D.E. Delorey, P.N. Pruneau — Boston College; AFCRL-TR-75-0303, December 1974.

HAES Report No. 22

Rocket-Borne Accelerator Module

D. Shephard, J. Carpenter, W. Reidy, W. Sheean, T. Zehnpfennig — Visidyne; VI-289, AFCRL-TR-75-0379, July 1975.

HAES Report No. 23

Instrumentation Analysis and Data Processing for Rocketburne LWIR Spectrometers (with Application to Rocket A18.006-2 of 22 March 1973)

J.W. Rogers - AFCRL; AFCRL-TR-75-0535, December 1975.

HAES Report No. 24

Auroral Simulation Studies in Support of ICECAP and EXCEDE

D. Archer, P. Tarr - Mission Research Corp; MRC-R-211, September 1975.

HAES Report No. 25*

Studies of Disturbed Ionospheres

K. Baker - Utah State; AFCRL-TR-75-0342, June 1975.

HAES Report No. 26°

Sky Radiance Calculations in the 0.5 µm - 5.0 µm Wavelength Range

W.G.M. Blättner, M.B. Wells — Radiation Research Associates; AFCRL-TR-75-0317, RRA-T7501, May 1975.

HAES Report No. 27

Data Reduction and Auroral Characterizations for ICECAP II

I.L. Kofsky, R.B. Sluder, C.A. Trowbridge - Photometrics; PhM 05-76, October 1975.

HAES Report No. 28°

Effective Recombination Coefficient of the Polar D Region Under Conditions of Intense Ionizing Radiation

T.M. Watt - SRI; DNA 3663T, July 1975.

HAES Report No. 29

Ionospheric Irregularities: HAES Program Support

R.D. Sears - Lockheed; DNA 3782F, September 1975.

^{*}HAES number omitted from this report. If you are a recipient of this report, please add the appropriate HAES report number to the front cover.

HAES Report No. 30°

Development of a Liquid-Helium Cooled Rocketborne Spectrometer

C.L. Wyatt, D.J. Baker — Utah State University; AFCRL-TR-75-0164, February 1975.

HAES Report No. 31*

SIMS Interferometer Study

R.W. Esplin, D.J. Baker, R.J. Huppi — Utah State University; AFCRL-TR-75-0500, July 1975.

HAES Report No. 32

Rocket Spectral Measurement of Atmospheric Infrared Emission During a Quiet Condition in the Auroral Zone

N.B. Wheeler and A.T. Stair, Jr. — AFGL; G. Frodsham and D.J. Baker — Utah State University; AFGL-TR-76-0252, October 1976.

HAES Report No. 33*

Report on the Geophysical Description and the Available Data Associated with Rocket PF-BB-53

G.J. Romick — Geophysical Institute, University of Alaska; AFCRL-TR-75-0040, January 1975.

HAES Report No. 34°

Report on the Geophysical Description and the Available Data Associated with Rocket PF-CI-97

G.J. Romick, Geophysical Institute, University of Alaska; AFCRL-TR-75-0327, May 1975.

HAES Report No. 35

A Correlation of Discrete and Diffuse Aurora With Particle Precipitation

R.S. Caverly, G.J. Romick, and R.D. Sharp — Geophysical Institute, University of Aleska; AFCRL-TR-75-0508, August 1975.

HAES Report No. 36°

Report on the Geophysical Description and the Available Data Associated with Rocket PF-HJ-NJ-90

G.J. Romick, Geophysical Institute, University of Alaska; AFCRL-TR-75-0362, December 1975.

HAES Report No. 371

ICECAP 73A Partial Reflection Sounder Results

G. Falcon - ITS; DNA 3943F, November 1975.

^{*}MAES number omitted from this report. If you are a recipient of this report, please add the appropriate MAES report number to the front cover.

¹DNA 3943F was erroneously issued as HAES Report No. 17 when in actuality it should have read "HAES Report No. 37." Holders of this report are urged to correct their copies.

Rocket Measurement of OH Emission Profiles in the 1.56 and 1.99 μm Bands
W.F. Grieder, K.D. Baker — Utah State University and A.T. Stair, AFCRL; AFCRLTR-76-0057, January 1976.

HAES Report No. 39

ICECAP 74 Chatanika Radar Results

M. Baron, SRI; DNA 3871T, October 1975.

HAES Report No. 40

Analyses of High-Altitude Effects Simulation (HAES)

W.P. Reidy, T.C. Degges, W. Neal — Visidyne; VI-311, AFGL-TR-76-0039, February 1976.

HAES Report No. 41

Geometrical Aspects of Rocket Photometry

W.F. Grieder — Utah State University, L.A. Whelan — Logicon Corp.; AFCRL-TR-76-0046, February 1976.

HAES Report No. 42*

Report on the Geophysical Description and Available Data Associated with Rocket PF-PT-81

G.J. Romick, Geophysical Institute, University of Alaska; AFCRL-TR-74-0540, October 1974.

HAES Report No. 43

ICECAP '75 Chatanika Radar Results

T.M. Watt - SRI; DNA 4086F, August 1976.

HAES Report No. 44

Results of Lower Ionospheric Measurements using the Partial Reflection Sounder During ICECAP '74

G. Falcon - ITS; Report number not yet assigned.

HAES Report No. 45°

EXCEDE: SWIR EXPERIMENT — Quick Look Data Report of 28 February 1976 Launch

R.R. O'Neil, A.T. Stair, Jr., J.C. Ulwick, R. Narcisi — AFGL; D. Burt — Utah State University; 3 March 1976.

HAES Report No. 46°

HIRIS EXPERIMENT - Quick Look Data Report of 1 April 1976 Launch

A.T. Stair, Jr. and J.W. Rogers — AFGL; W.R. Williamson-Honeywell Radiation Center; AFGL-OP-TM-02, July 1976.

^{*}HAES number omitted from this report. If you are a recipient of this report, please add the appropriate HAES report number to the front cover.

HAES Report No. 47°

ICECAP '74 - Chatanika Radar Results

P.D. Perreault, M.J. Baron - SRI; DNA 3871T, October 1975.

HAES Report No. 48°

Bias and Signal Processing Circuits for a Mass Spectrometer in the Project EXCEDE: SWIR Experiment

R. Sukys, J.S. Rochefort, S. Goldberg — Northeastern University; AFGL-TR-76-0060, October 1975.

HAES Report No. 49

Chatanika Radar Results during the EXCEDE Experiment

T.M. Watt - SRI; DNA 4123T, September 1976.

HAES Report No. 50

Rocketborne Measurements of Infrared Enhancements Associated with a Bright Auroral Breakup

K.D. Baker, D.J. Baker — Utah State University, J.C. Ulwick, A.T. Stair — AFGL. Report number not yet assigned.

HAES Report No. 51

LWIR (7-24 µm) Measurements from the Launch of a Rocketborne Spectrometer into an Aurora (1973)

J.W. Rogers, A.T. Stair, Jr., N.B. Wheeler — AFGL; C.L. Wyatt, D.J. Baker — Utah State University. AFGL - TR - 76 - 0274.

HAES Report No. 52*

DNA Project 609 Radar: Auroral Backscatter Measurements

R.T. Tsunoda, R.I. Presnell, T.N.C. Wang - SRI; DNA 3929F, February 1976.

HAES Report No. 53*

ICECAP Data Processing System

D.E. Delorey and P.N. Pruneau -- Boston College; AFGL-TR-76-0138, May 1976.

^{*}HAES number omitted from this report. If you are a recipient of this report, please add the appropriate HAES report number to the front cover.

Appendix C

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Distribution List

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OTCY	ATTH CODE 7700 TIMOTHY P COFFEY
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